

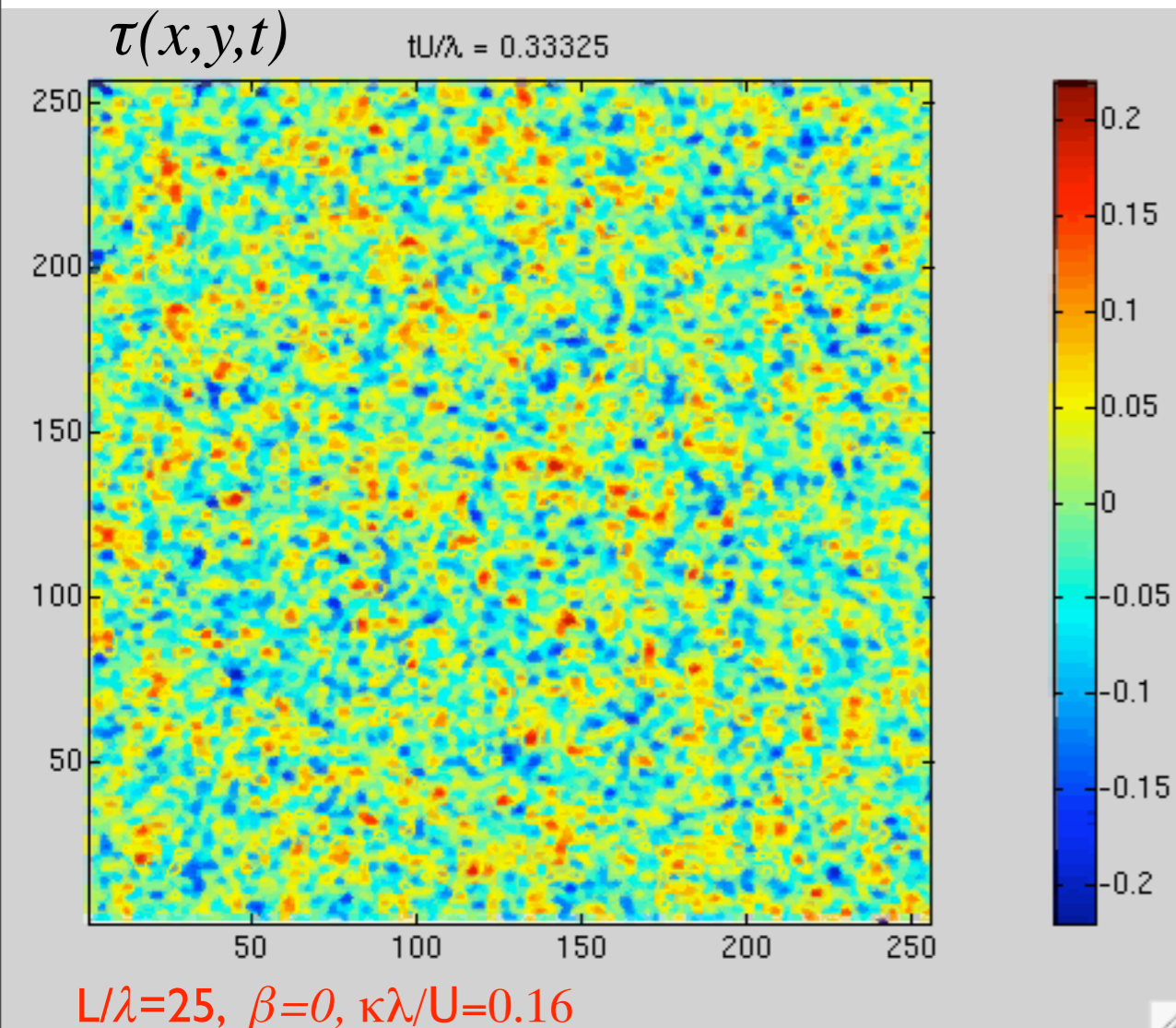


Bottom Drag, eddy diffusivity, wind work
and the power integrals

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Two-layer baroclinic turbulence



The eddy diffusivity of heat is:

$$D_\tau \equiv U^{-1} \langle \psi_x \tau \rangle$$

above $\langle \rangle = x, y, t$ average

Dimensional considerations (HH80) show:

$$D_\tau = U\lambda \times D_\tau^* \left(\frac{\lambda}{L}, \frac{\kappa\lambda}{U}, \frac{\beta\lambda^2}{U}, \frac{\nu}{UL^7} \right)$$

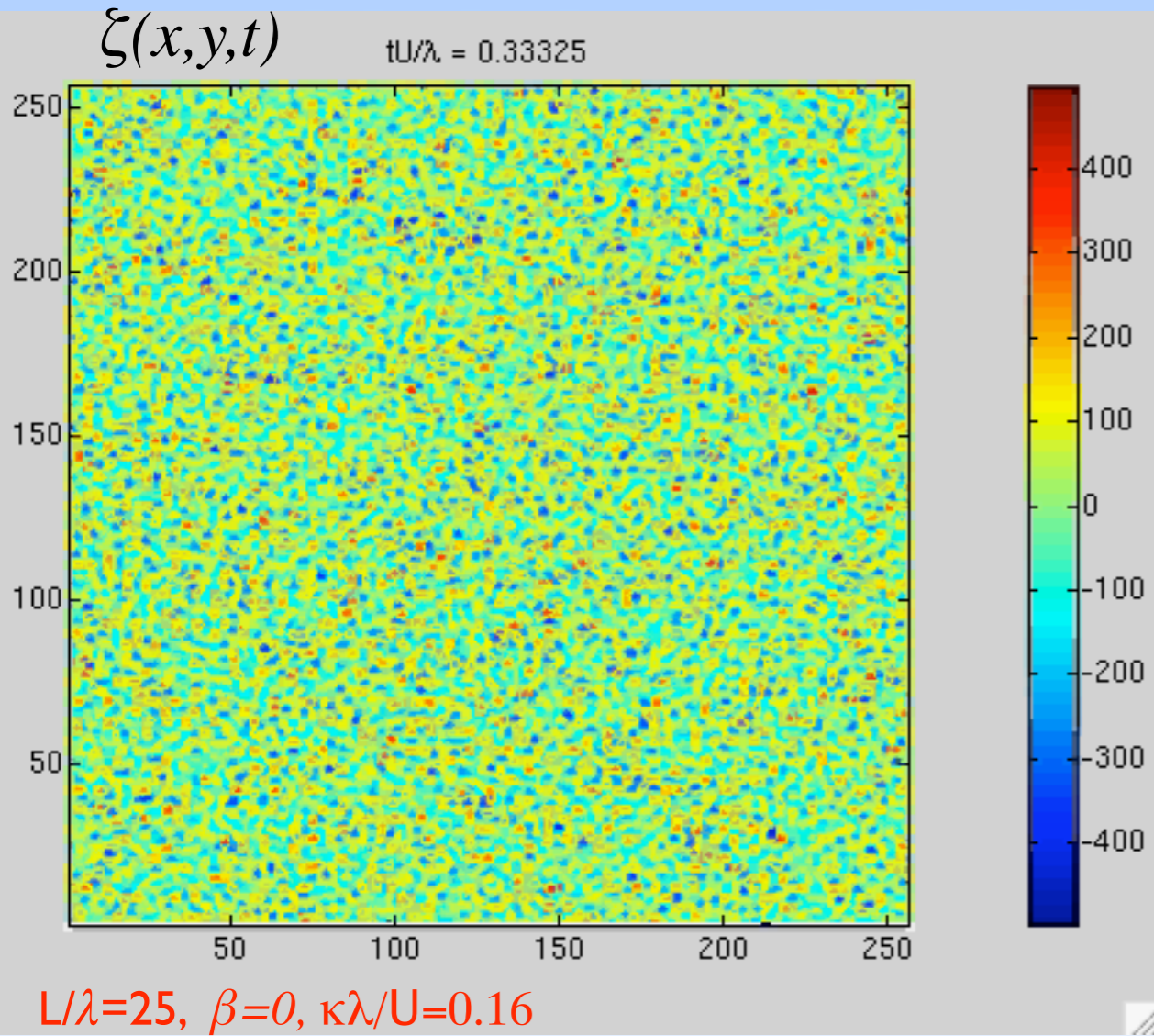
This is the cleanest example of the eddy diffusivity problem.

$$\nabla^2 \psi_t + J(\psi, \nabla^2 \psi) + J(\tau, \nabla^2 \tau) + U \nabla^2 \tau_x = -\kappa \nabla^2 (\psi - \sqrt{2} \tau) - \nu \nabla^8 (\nabla^2 \psi),$$

$$\begin{aligned} (\nabla^2 - \lambda^{-2}) \tau_t + J(\psi, (\nabla^2 - \lambda^{-2}) \tau) + J(\tau, \nabla^2 \psi) + U (\nabla^2 + \lambda^{-2}) \psi_x \\ = -\kappa \nabla^2 (2\tau - \sqrt{2} \psi) - \nu \nabla^8 (\nabla^2 - \lambda^2) \tau. \end{aligned}$$

The eddy diffusivity of heat is:

$$D_\tau \equiv U^{-1} \langle \psi_x \tau \rangle$$



Recent work on parameterizing D (LH95, HL96, LH03) results in:

$$D_\tau^{LH03} = U \lambda \times 1.75 \beta_*^{-2} (1 - \beta_*)^{5/2}$$

These LH theories are based on inertial range arguments (R77, S78, S80), and ignore the **power integral**:

$$U^2 \lambda^{-2} D_\tau \approx \kappa \left\langle |\nabla \psi - \sqrt{2} \nabla \tau|^2 \right\rangle$$

$$\frac{1}{2} \frac{\partial}{\partial t} \left(\langle |\nabla \psi|^2 \rangle + \langle |\nabla \tau|^2 \rangle + \lambda^{-2} \langle \tau^2 \rangle \right) + U \lambda^{-2} \langle \psi_x \tau \rangle = \kappa \left\langle |\nabla \psi - \sqrt{2} \nabla \tau|^2 \right\rangle + \text{ssd}$$

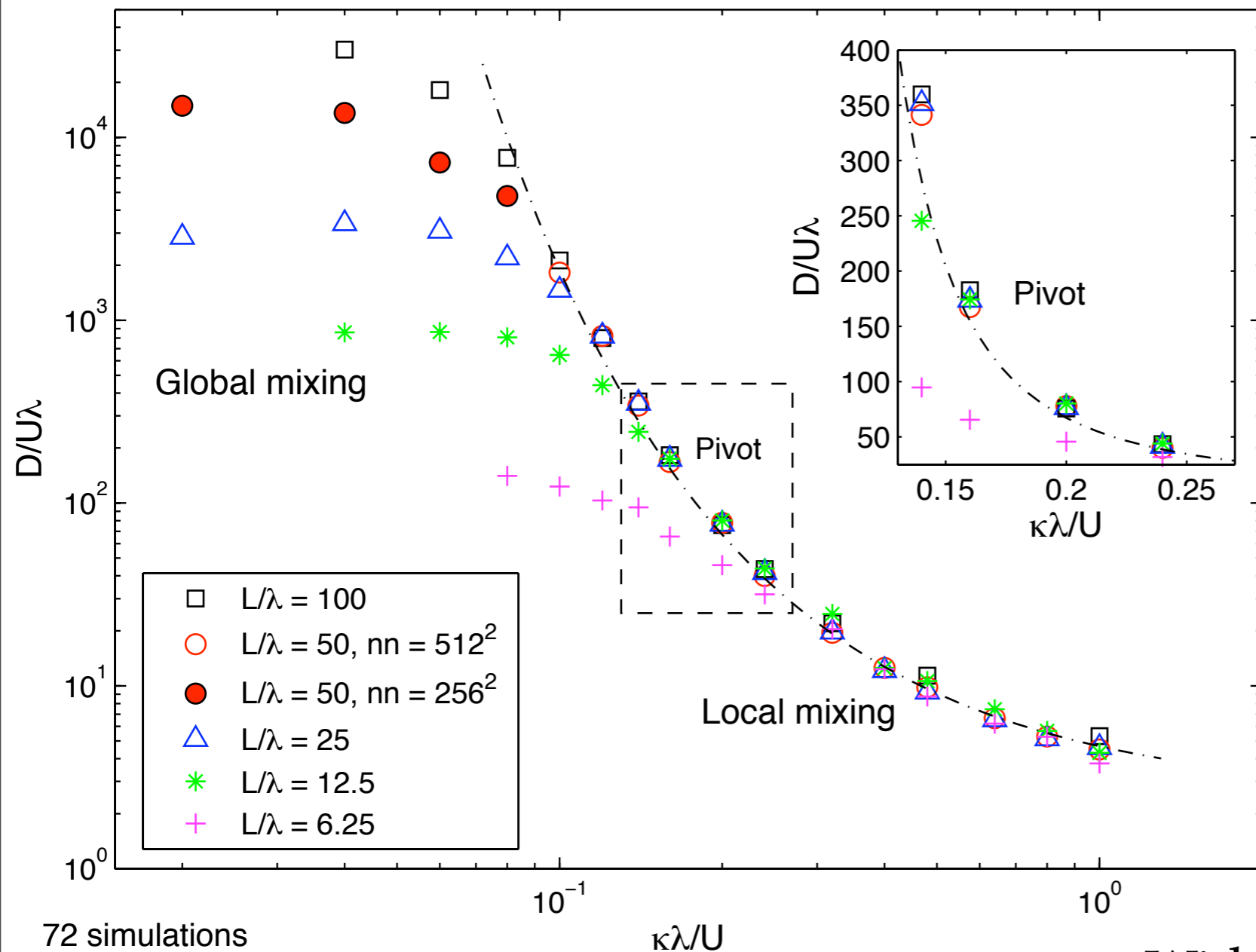
Time rate of change of energy

Energy
production

Ekman
dissipation

Hyper-
viscosity

D from 72 simulations with $\beta=0$



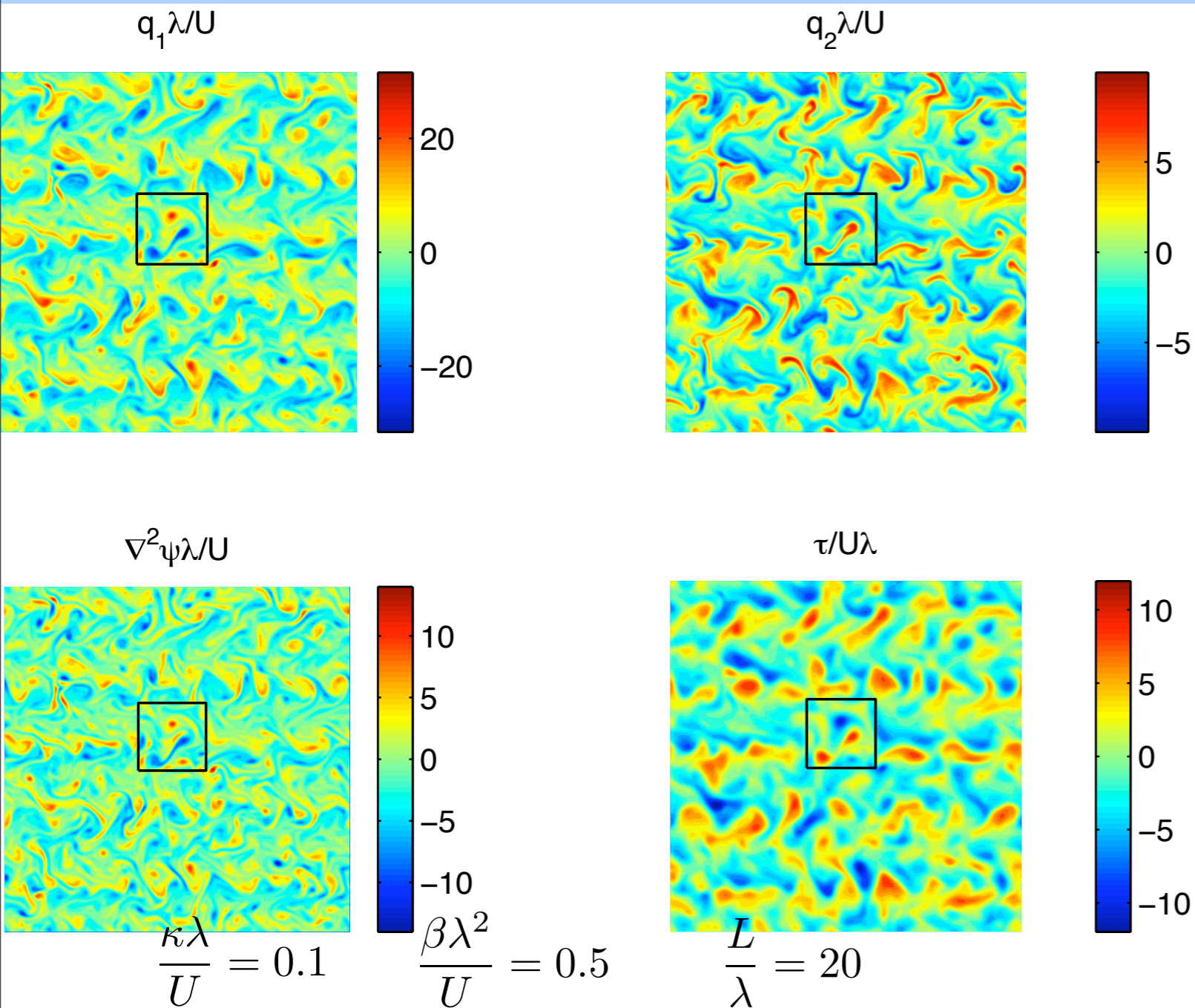
$$D = U\lambda \times D_* \left(\frac{L}{\lambda}, \frac{\kappa\lambda}{U}, 0 \right)$$

The dashed curve is:

$$D_* \approx 2.4 \exp(2U/3\kappa\lambda)$$

With $\beta=0$, D is **very** sensitive to bottom drag --- the inverse cascade is halted only by bottom drag.

Now the β -effect



The inverse cascade of the barotropic mode can be arrested by β (without dissipation).

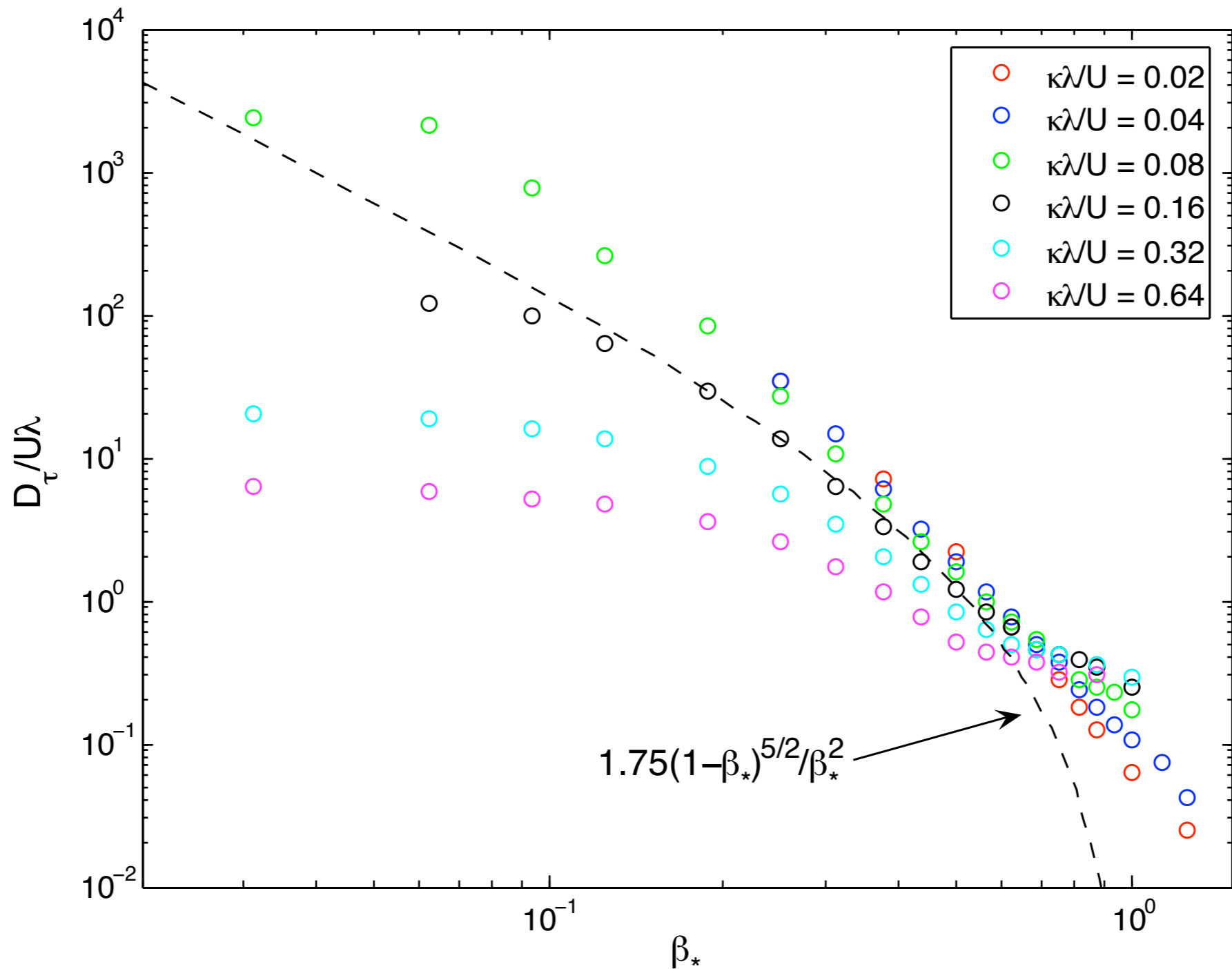
$$\ell_{\text{Rhines}} = \sqrt{\frac{U}{\beta}}$$

Perhaps non-dissipative scaling arguments work?

$$D_{\tau}^{LH03} = U \lambda \times 1.75 \beta_*^{-2} (1 - \beta_*)^{5/2}$$

However we still have: $U^2 \lambda^{-2} D_{\tau} \approx \kappa \left\langle |\nabla \psi - \sqrt{2} \nabla \tau|^2 \right\rangle$

QG eddy diffusivity, non-zero β



$L/\lambda=25$

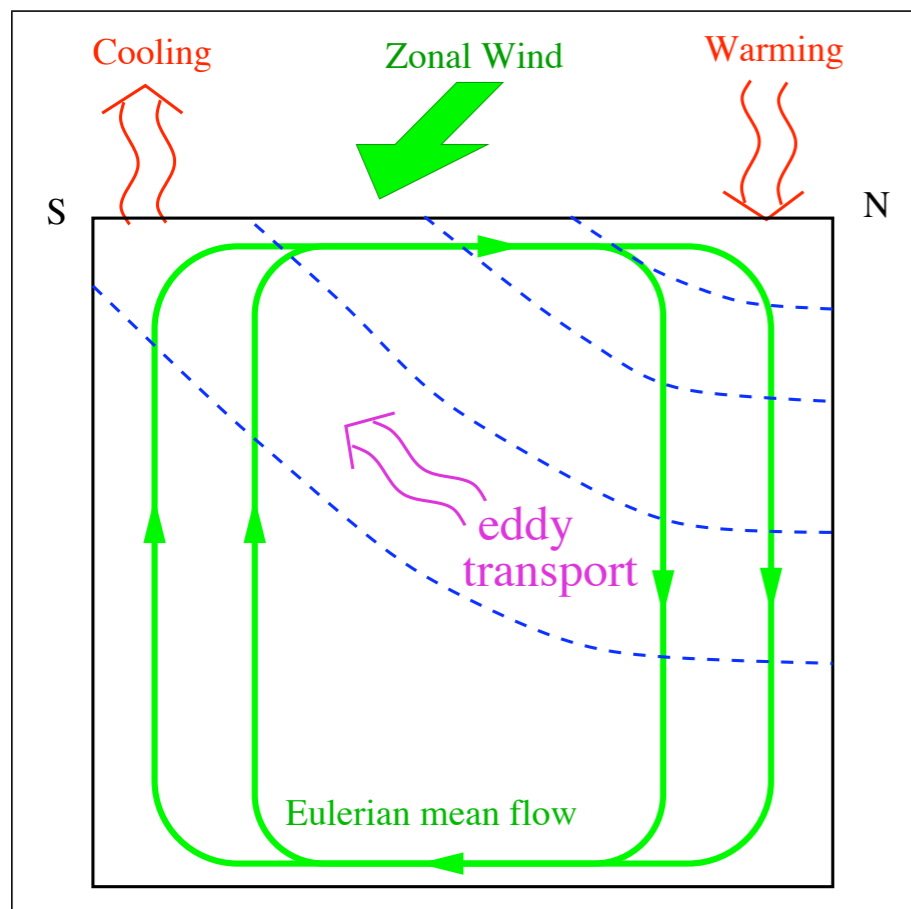
$$D_{\tau}^{LH03} = U\lambda \times 1.75\beta_*^{-2}(1-\beta_*)^{5/2}$$

QG conclusions

- There is **no** part of the parameter space where bottom drag is not qualitatively important.
- There are **two** decisive parameters, β and κ .
- The power integral is important e.g., the TY05 parametrization of D .
- Is there a useful generalization of the power integral to PE models?

$$U^2 \lambda^{-2} D_\tau \approx \kappa \left\langle |\nabla \psi - \sqrt{2} \nabla \tau|^2 \right\rangle$$

Idealized models of the ACC



from Karsten, Jones & Marshall (2002)

How does the thermocline depth, h , depend on external parameters? Using TEM arguments:

$$\overline{v'b'} \approx -D_{\text{eddy}} \bar{b}_y \approx \frac{\tau_s}{f_0} \bar{b}_z \quad \Rightarrow \quad h \sim \frac{\tau L_y}{f D_{\text{eddy}}}$$

How to find the eddy- D ? For example, KJM suggest:

$$D_{\text{eddy}}^{\text{KJM}} = L_y \times \frac{g'h}{f L_y}$$

$$\frac{D\mathbf{u}}{Dt} + \hat{\mathbf{z}} \times f_0 \mathbf{u} + \nabla p = \nabla \cdot \nu \nabla \mathbf{u} + b \hat{\mathbf{z}} + \gamma_s(z) \tau_s(z) \hat{\mathbf{x}} - r \gamma_b(z) \mathbf{u} ,$$

$$\frac{Db}{Dt} = \nabla \cdot \kappa \nabla b ,$$

$$\nabla \cdot \mathbf{u} = 0 .$$

Bottom drag is now r , and the wind stress is τ .

The zonal mean bottom velocity is determined by r . $\bar{u}_b(y) = \frac{\tau_s(y)}{rH}$

The power integral

In this problem the power integral is:

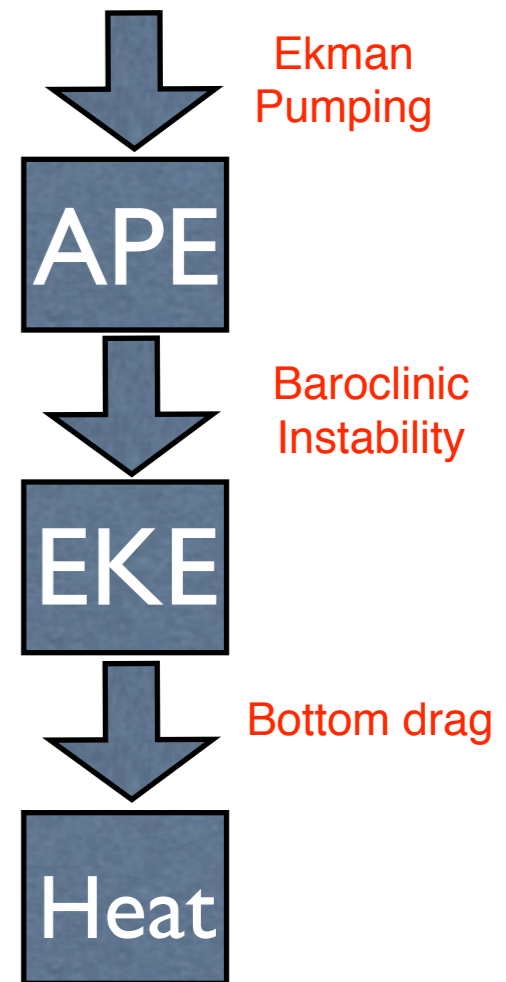
$$\langle w'b' \rangle \underset{\text{PY05}}{\approx} - \langle \bar{w}\bar{b} \rangle \underset{\text{GG574}}{\approx} r \langle |\mathbf{u}'_b|^2 \rangle$$

TEM arguments give:

$$\langle w'b' \rangle \approx - \left\langle v'b' \frac{\bar{b}_y}{\bar{b}_z} \right\rangle \approx \left\langle D_{\text{eddy}} \frac{\bar{b}_y^2}{\bar{b}_z} \right\rangle$$

Note: $\bar{w} \approx w_E = -\frac{1}{f_0} \frac{d\tau_s}{dy}$ and so: $-\langle \bar{w}\bar{b} \rangle \sim \frac{h}{H} \frac{\tau g'}{f_0 L_y} \sim r u_b'^2$

CPY05 “eddy transfer velocity”.



The power integral - derive GGS74

$$\underbrace{\langle \gamma_s \tau_s u \rangle}_{\text{Wind work}} + \underbrace{\langle wb \rangle}_{\text{PE} \rightarrow \text{KE}} = r \underbrace{\langle \gamma_b |\mathbf{u}|^2 \rangle}_{\text{Bottom drag}} + \underbrace{\langle \nu \|\nabla \mathbf{u}\|^2 \rangle}_{\text{viscosity}}$$

After some simplifications: $H^{-1} \langle \tau_s \bar{u}_s \rangle \approx r \langle |\mathbf{u}_b|^2 \rangle = r \langle \bar{u}_b^2 \rangle + r \langle |\mathbf{u}'_b|^2 \rangle$
 (N.B. true even with topography)

How about PE \rightarrow KE??? $\langle wb \rangle = \langle \kappa b_z \rangle$
↑
very small

Now estimate the wind work using: $\bar{u}_s \approx \bar{u}_b - \frac{1}{f_0} \int_{-H}^0 \bar{b}_y(y, z) dz$

Thus: $-H^{-1} \left\langle \int_{-H}^0 \bar{b}(y, z) dz w_E(y) \right\rangle \approx r \langle |\mathbf{u}'_b|^2 \rangle$

Note: $\bar{w} \approx w_E = -\frac{1}{f_0} \frac{d\tau_s}{dy}$ and so: $\frac{h}{H} \frac{\tau g'}{f_0 L_y} \sim r u_b^2$