

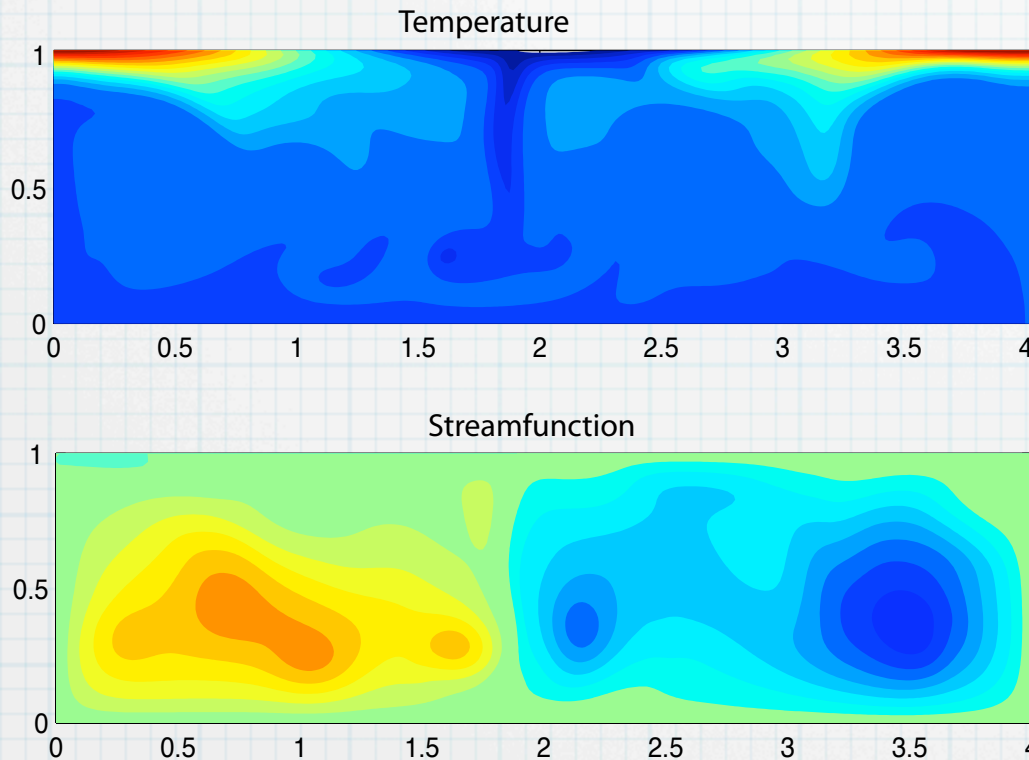
# Sandström's theorem and the energetics of ocean mixing

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Sandström's (1908) "theorem": a closed steady circulation can only be maintained if the heat source is beneath the cold source (Defant 1965).



A 2-D counterexample: the box is heated non-uniformly at the top.

This is Rossby's horizontal convection (HC) problem.

The critical Rayleigh number of HC is zero --- the slightest  $\Delta T$  imposed at the top surface causes motion.

# Why bother with Sandström?

- \* There are many recent citations and endorsements of the 1908 “theorem”: Munk & Wunsch (1998), Huang (1999), Emmanuel (2001), Wunsch & Ferrari (2004), etc.
- \* And also many counterexamples against the “theorem”: Jeffreys (1925), Rossby (1965), Mullaney et al. (2004), Wang & Huang (2005), etc.
- \* What controls the THC? What drives (i.e., supplies energy to) the THC?

Sandström's (1908) “theorem”: a closed steady circulation can only be maintained if the heat source is beneath the cold source (from Defant 1965).

# Why isn't HC universally accepted as a counterexample to the "theorem"?

- \* HC isn't a vigorous flow.
- \* HC produces a thin thermocline --- you need wind, tides and breaking IGW's to explain deep ocean stratification.
- \* "strict interpretation of the theorem is difficult" (Houghton 1997).

The goal here: make the "theorem" into a theorem.



An enclosed Boussinesq fluid --- like the ocean.



$$\mathbf{u} \cdot \hat{\mathbf{n}} = 0$$

$$\frac{D\mathbf{u}}{Dt} + \hat{\mathbf{z}} \times f\mathbf{u} + \nabla p = g\alpha T\hat{\mathbf{z}} + \nu\nabla^2\mathbf{u}$$

$$\nabla \cdot \mathbf{u} = 0$$

$$\frac{DT}{Dt} = \kappa\nabla^2 T$$

Note: linear EOS

# Notation

- = area and time average,
- ⟨•⟩ = volume and time average

For example:

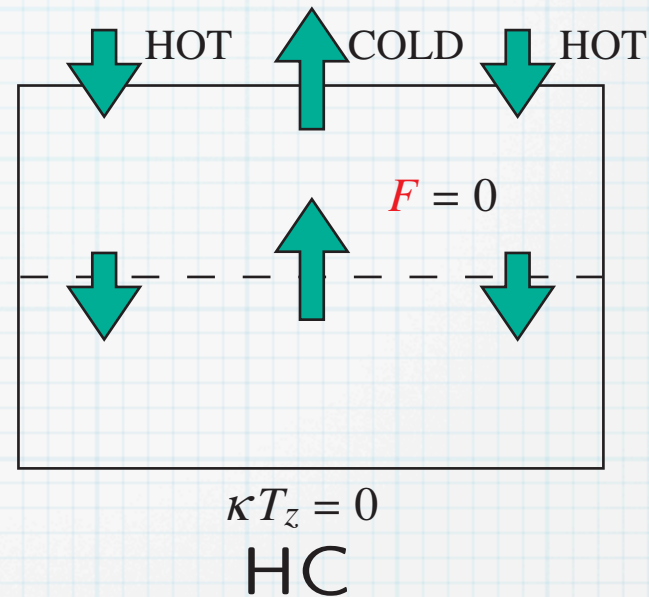
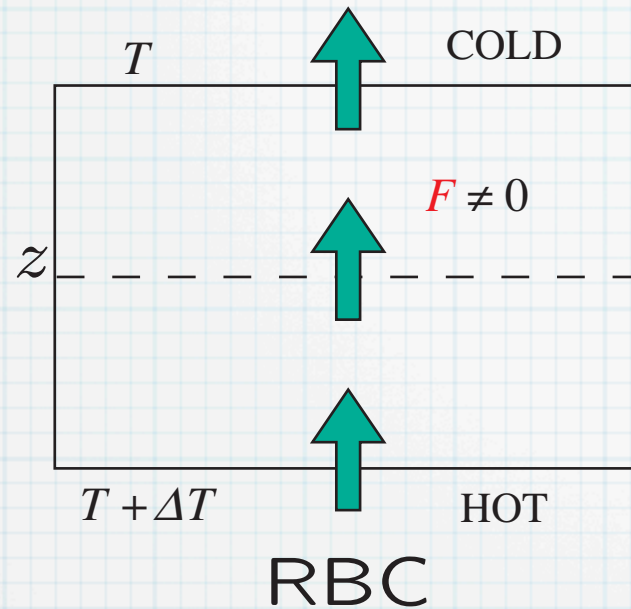
$$\overline{wT} = A^{-1} t_{\infty}^{-1} \int_0^{t_{\infty}} \iint wT \, dx dy \, dt$$

$$KE = \langle \frac{1}{2} |\mathbf{u}|^2 \rangle = V^{-1} t_{\infty}^{-1} \int_0^{t_{\infty}} \iiint \frac{1}{2} |\mathbf{u}|^2 \, dx dy dz \, dt$$

where  $V = H \times A$ .

We assume that all these averages are stationary.

# The flux constraint



Average over  $x, y$  and  $t$  :

$$\frac{DT}{Dt} = \kappa \nabla^2 T \quad \Rightarrow \quad \overline{wT} - \kappa \bar{T}_z = F \quad (\forall z)$$

$F$  = the vertical heat flux

The momentum equation has not been used.

To obtain the **the mechanical energy budget**, take:

$$\int \mathbf{u} \cdot \left[ \frac{D\mathbf{u}}{Dt} + \hat{\mathbf{z}} \times f\mathbf{u} + \nabla p = g\alpha T\hat{\mathbf{z}} + \nu\nabla^2\mathbf{u} \right] dV$$

After  $\langle \rangle$ , viscous dissipation = PE release:

$$\underbrace{\nu \int |\nabla\mathbf{u}|^2 dV}_{\varepsilon} = g\alpha \langle wT \rangle$$

Now recall the zero-flux constraint:

$$\overline{wT} - \kappa\bar{T}_z = 0 \quad \Rightarrow \quad \langle wT \rangle = \frac{\kappa\Delta\bar{T}}{H}$$

Eliminate  $\langle wT \rangle$ :

$$\boxed{\underbrace{\nu \langle |\nabla\mathbf{u}|^2 \rangle}_{\varepsilon} = \kappa \frac{\alpha g \Delta\bar{T}}{H}}$$

(A rehabilitation of Sandström's theorem.)

The top-to-bottom  $\Delta\bar{T}$  is:  $\Delta\bar{T} = \bar{T}(0) - \bar{T}(-H)$



An estimate using:

$$\underbrace{\nu \langle |\nabla \mathbf{u}|^2 \rangle}_{\varepsilon} = \kappa \frac{\alpha g \Delta \bar{T}}{H}$$

We can easily estimate the mechanical energy dissipation:

$$\varepsilon \approx 10^{-7} \times \frac{2 \times 10^{-2}}{4000} = 2 \times 10^{-12} \text{ W kg}^{-1}.$$

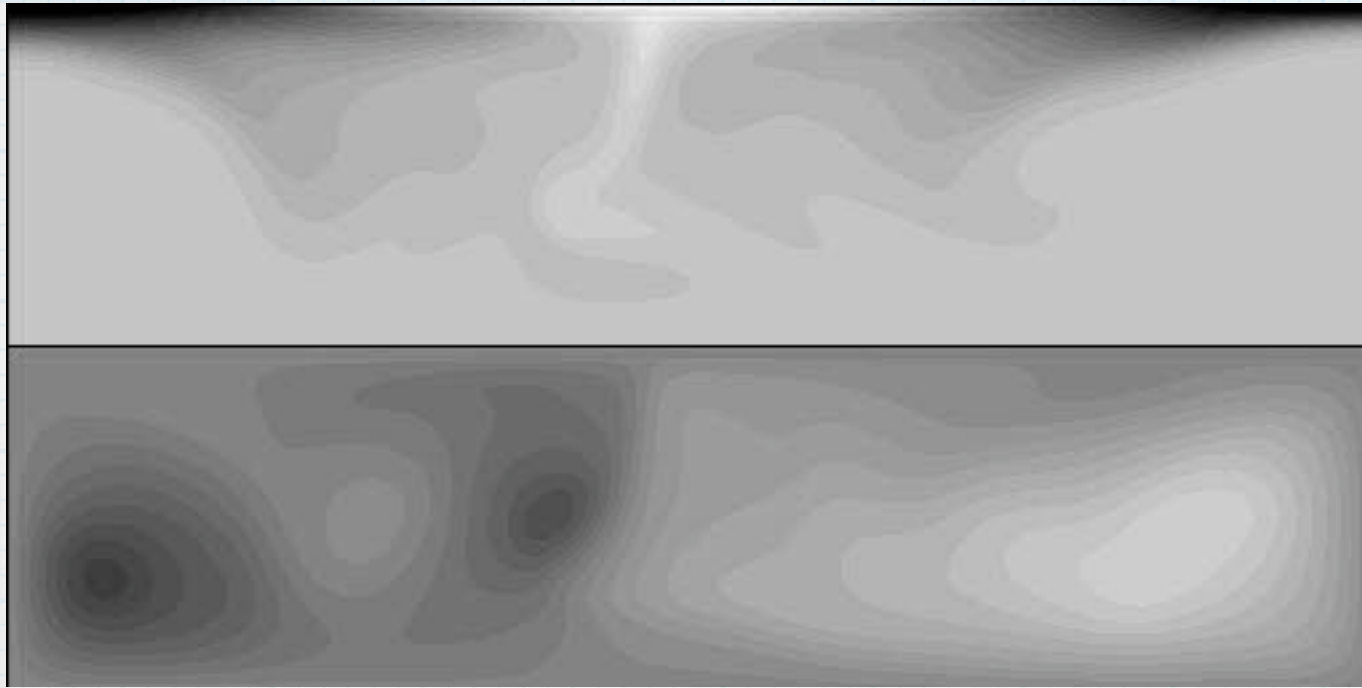
This is smaller by a factor of one thousand than observations.

**Conclusion 1:** the observed level of ocean turbulence requires other sources of KE (winds, tides and fish).

**Conclusion 2:** there is a kernel of truth in Sandström's theorem, or:

**PY (2002) theorem:** the mechanical energy supplied by non-uniformly heating only the top surface is directly proportional to the **molecular** diffusion of heat.

# Back to the zero-flux constraint



At every depth:

$$\overline{wT} - \kappa \overline{T}_z = 0$$

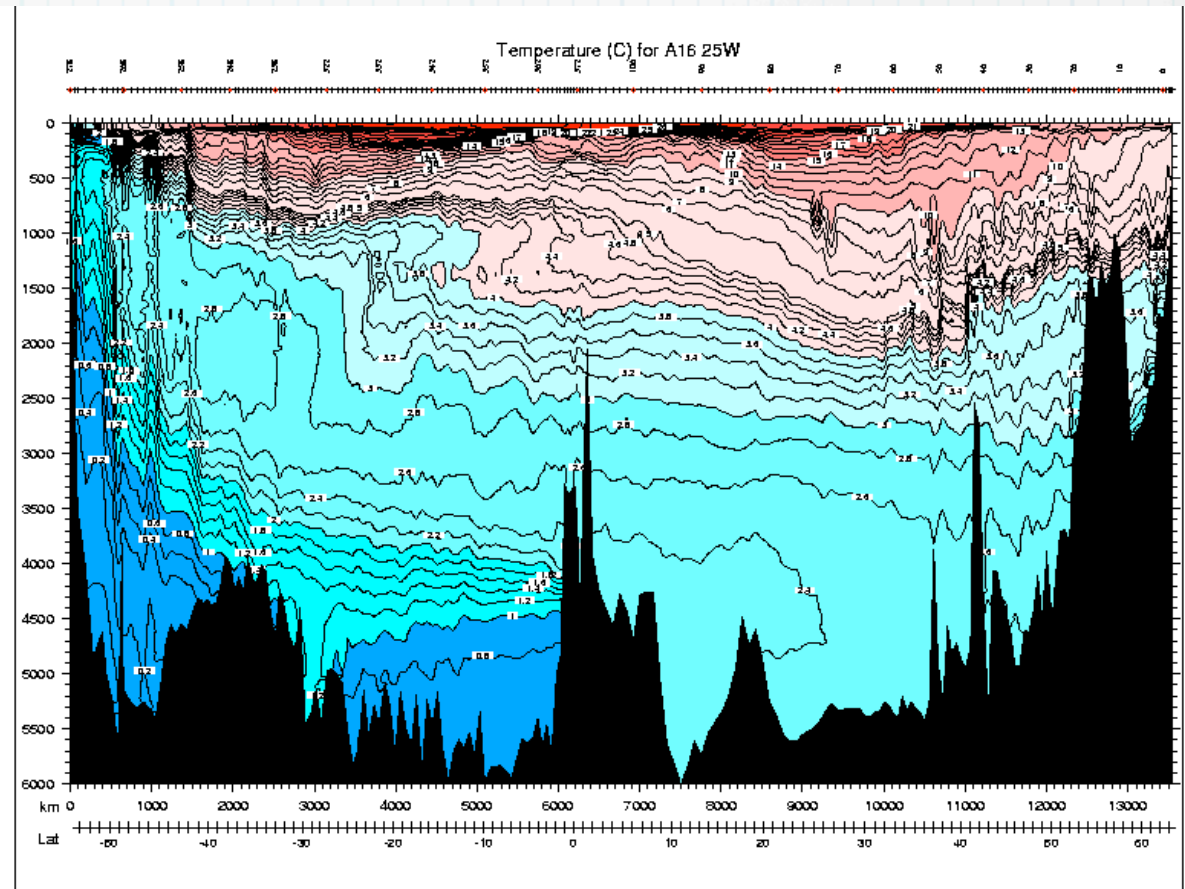
Although the energy supply is small, the flow is **not laminar**.

The zero-flux constraint follows only from the thermodynamic equation and the bottom BC --- no consideration of the momentum equation is necessary.

# THE OCEAN....

- \* The zero flux constraint almost\* applies.
- \* There is almost\* no net conversion between PE and KE in the ocean.

\* "almost" because of molecular diffusion, geothermal heating, nonlinear EOS and radiation



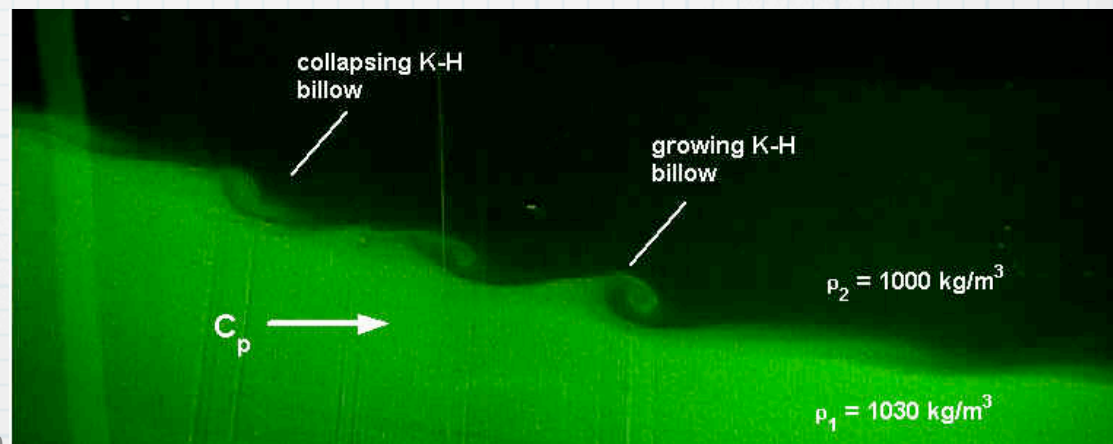
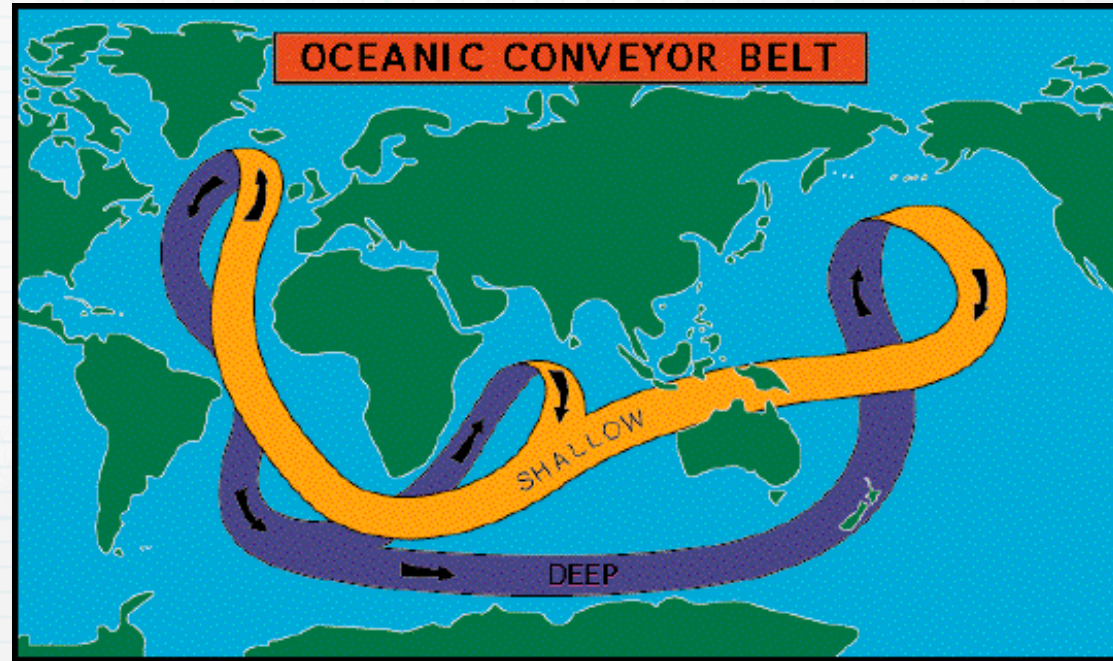
$$\overline{wT} \approx 0 \quad \nabla z$$



$$\langle wT \rangle = \underbrace{\langle \bar{w}\bar{T} \rangle}_{\text{large scales} > 0} + \underbrace{\langle w'T' \rangle}_{\text{small scales} < 0} \approx 0$$

- \* Large-scale THC lowers the center of gravity.
- \* Small-scale mixing raises the center of gravity.
- \* These processes are almost\* in perfect balance.
- \* Large scale overturning is linked to small scale mixing.

\* "almost" because of molecular diffusion, geothermal heating, nonlinear EOS and radiation



EFML, Stanford

TMFKAWS, June 2005



# Conclusions

- \* Sandström's arguments are not completely wrong...
- \* The mechanical energy supplied by non-uniform surface heating is directly proportional to **molecular** diffusion.
- \* The net vertical buoyancy flux in the ocean is a small energy source (e.g., less than bioturbation).
- \* The zero-flux constraint links small-scale mixing and large-scale TH overturning.