

# Non-geostrophic Baroclinic Stability

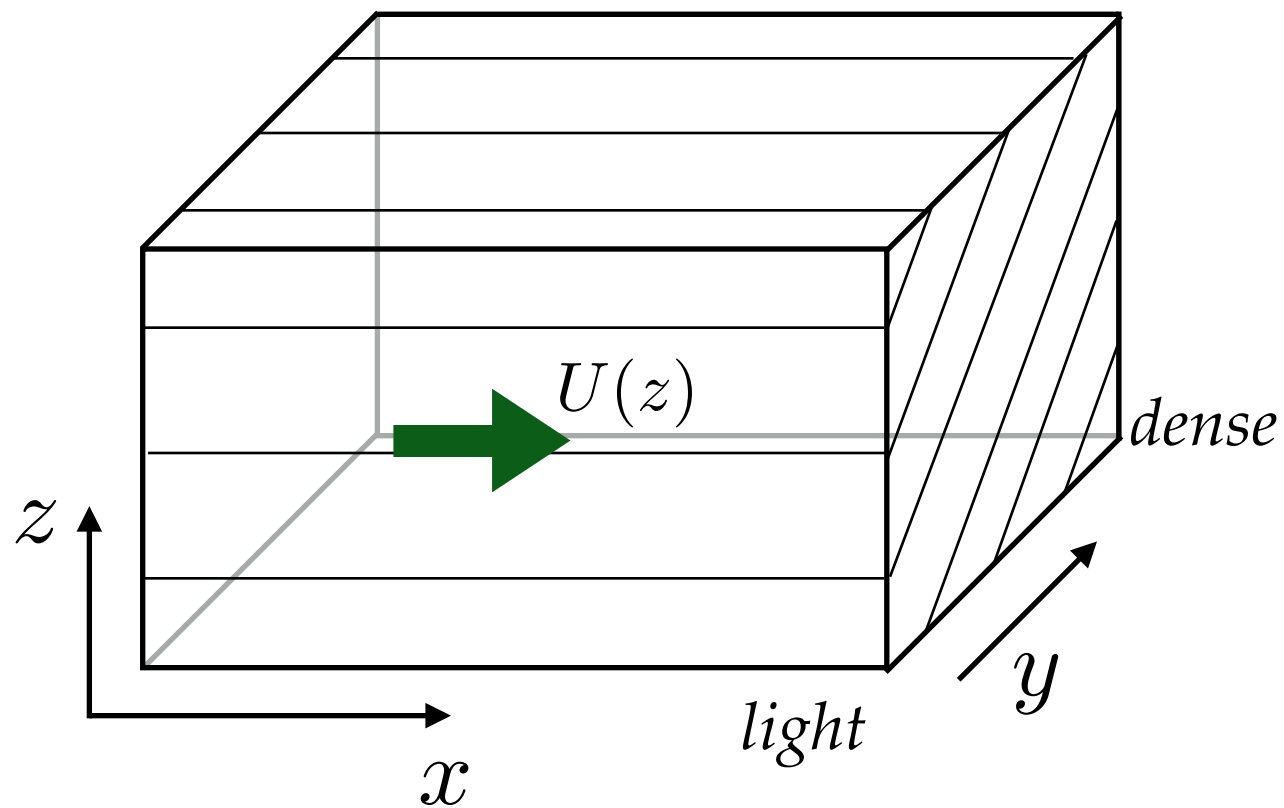
Stone, PH. *On Non-Geostrophic Baroclinic Stability*, JAS 1966

Stone, PH. *On Non-Geostrophic Baroclinic Stability: Part II*, JAS 1966

Haine, TWN. & Marshall, J. *Gravitational, Symmetric and Baroclinic Instability of the Ocean Mixed Layer*, JPO 1998

Holton, *An Introduction to Dynamic Meteorology. Fourth Edition.*

# The background state is in thermal wind balance



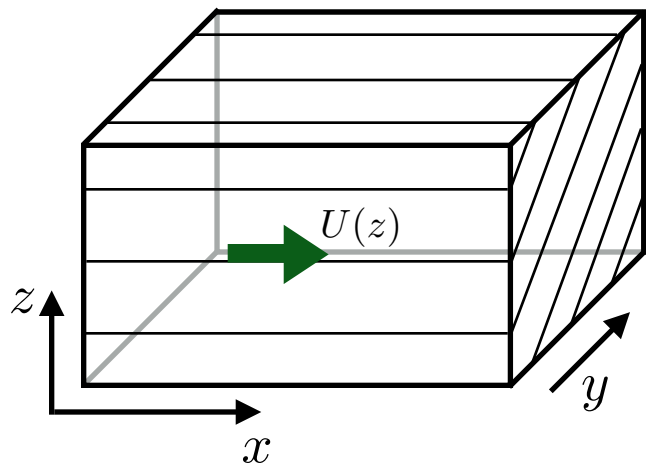
$$N^2 = -\frac{g}{\rho_0} \frac{\partial \bar{\rho}}{\partial z} = \bar{b}_z$$

$$M^2 = -\frac{g}{\rho_0} \frac{\partial \bar{\rho}}{\partial y} = \bar{b}_y$$

$$f \frac{dU}{dz} = \frac{g}{\rho_0} \frac{\partial \bar{\rho}}{\partial y} = -M^2$$

$$U = -\frac{M^2 z}{f}$$

# Primitive equations (not QG!), linearized around the background state



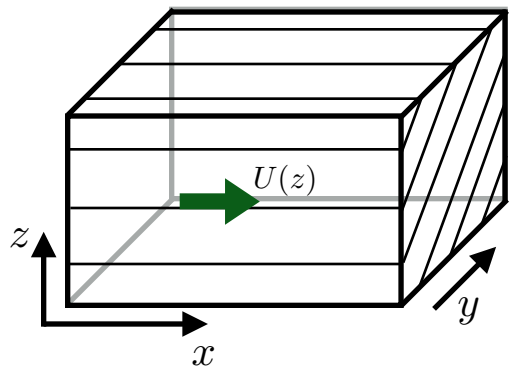
Background + perturbations:

$$\begin{aligned}u' &= u + U & v' &= v & w' &= w \\b' &= b + \bar{b} & p' &= p + \bar{p}\end{aligned}$$

Linearized set of equations:

$$\begin{aligned}u_t + Uu_x + wU' - fv &= -p_x \\v_t + Uv_x + fu &= -p_y \\0 &= -p_z + b \\b_t + Ub_x + v\bar{b}_y + w\bar{b}_z &= 0 \\u_x + v_y + w_z &= 0\end{aligned}$$

# Primitive equations (not QG!), linearized around the background state



non-dimensionalize:

$$\begin{aligned}
 (x^*, y^*) &= \frac{U}{f}(x, y) & z^* &= H_0 z \\
 (u^*, v^*) &= U(u, v) & w^* &= H_0 f w \\
 b^* &= N^2 H_0 b & p^* &= N^2 H_0^2 p \\
 t^* &= \frac{1}{f} t & \sigma^* &= f \sigma \\
 H^* &= \frac{1}{H_0} H \\
 & e^{i(kx + ly + \sigma t)}
 \end{aligned}$$

Assume a solution like:

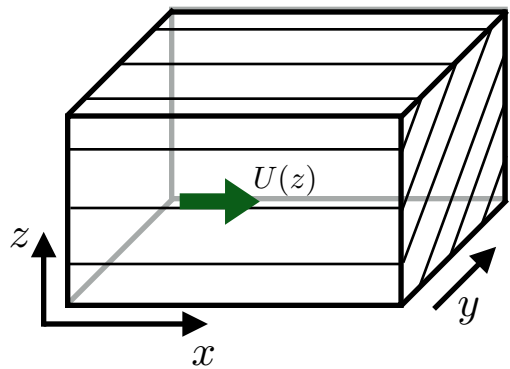
Combine to form eigenvalue equation:

$$[1 - (\sigma + kU(z))^2]w'' - 2 \left[ \frac{k}{\sigma + kU(z)} - il \right] w' - \left[ Ri(k^2 + l^2) - \frac{2ikl}{\sigma + kU(z)} \right] w = 0$$

Boundary conditions  $w(0) = w(1) = 0$

$$Ri = \frac{N^2 H_0^2}{U^2} = \frac{N^2 f^2}{M^4}$$

# Instabilities of the non-geostrophic eigenvalue equation

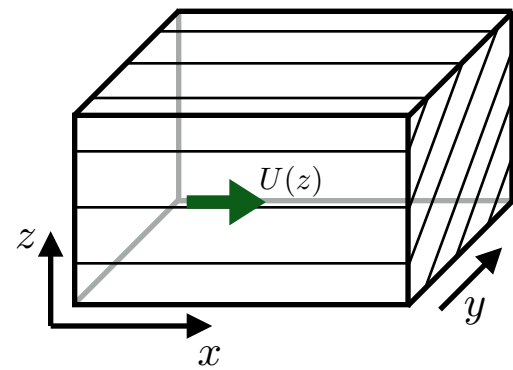


$$[1 - (\sigma + kU(z))^2]w'' - 2 \left[ \frac{k}{\sigma + kU(z)} - il \right] w' - \left[ Ri(k^2 + l^2) - \frac{2ikl}{\sigma + kU(z)} \right] w = 0$$

Boundary conditions  $w(0) = w(1) = 0$

- Gravitational ( $Ri < 0$ )
- Kelvin-helmholtz ( $0 < Ri < 0.25$ )
- Symmetric ( $k=0$ ,  $Ri < 1$ )
- Baroclinic ( $l=0$ ,  $Ri > 1$ )
- Also allows inertia-gravity waves!

# No zonal derivatives: symmetric instability



Let  $k = 0$

Eigenvalue equation:  $[1 - \sigma^2]w'' + 2ilw' - \text{Ri} l^2 w = 0$

Assume solution:  $w = w_0 e^{imz}$

$$m^2 + \frac{2l}{1 - \sigma^2} m + \frac{\text{Ri} l^2}{1 - \sigma^2} = 0$$

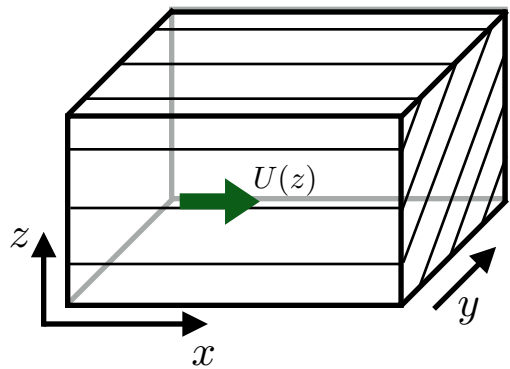
$$m = \frac{l}{\sigma^2 - 1} \pm \sqrt{\frac{l^2}{(\sigma^2 - 1)^2} + \frac{\text{Ri} l^2}{\sigma^2 - 1}}$$

$$w = w_0 (a e^{im_+ z} + b e^{im_- z})$$

Boundary conditions:  $w = 0$  at  $z = 0, 1$

$$\implies m_+ - m_- = 2n\pi \quad \text{and} \quad a = -b$$

# Symmetric instability: growth rates



Boundary conditions impose

$$n\pi = \sqrt{\frac{l^2}{(\sigma^2 - 1)^2} + \frac{\text{Ri}l^2}{\sigma^2 - 1}}$$

Growth rate: 
$$\sigma^2 = 1 + \frac{\text{Ri}l^2}{2n^2\pi^2} \pm \sqrt{\frac{\text{Ri}^2l^4}{4n^4\pi^4} + \frac{l^2}{n^2\pi^2}}$$

With our assumed solution:  $e^{i(kx+ly+\sigma t)}$

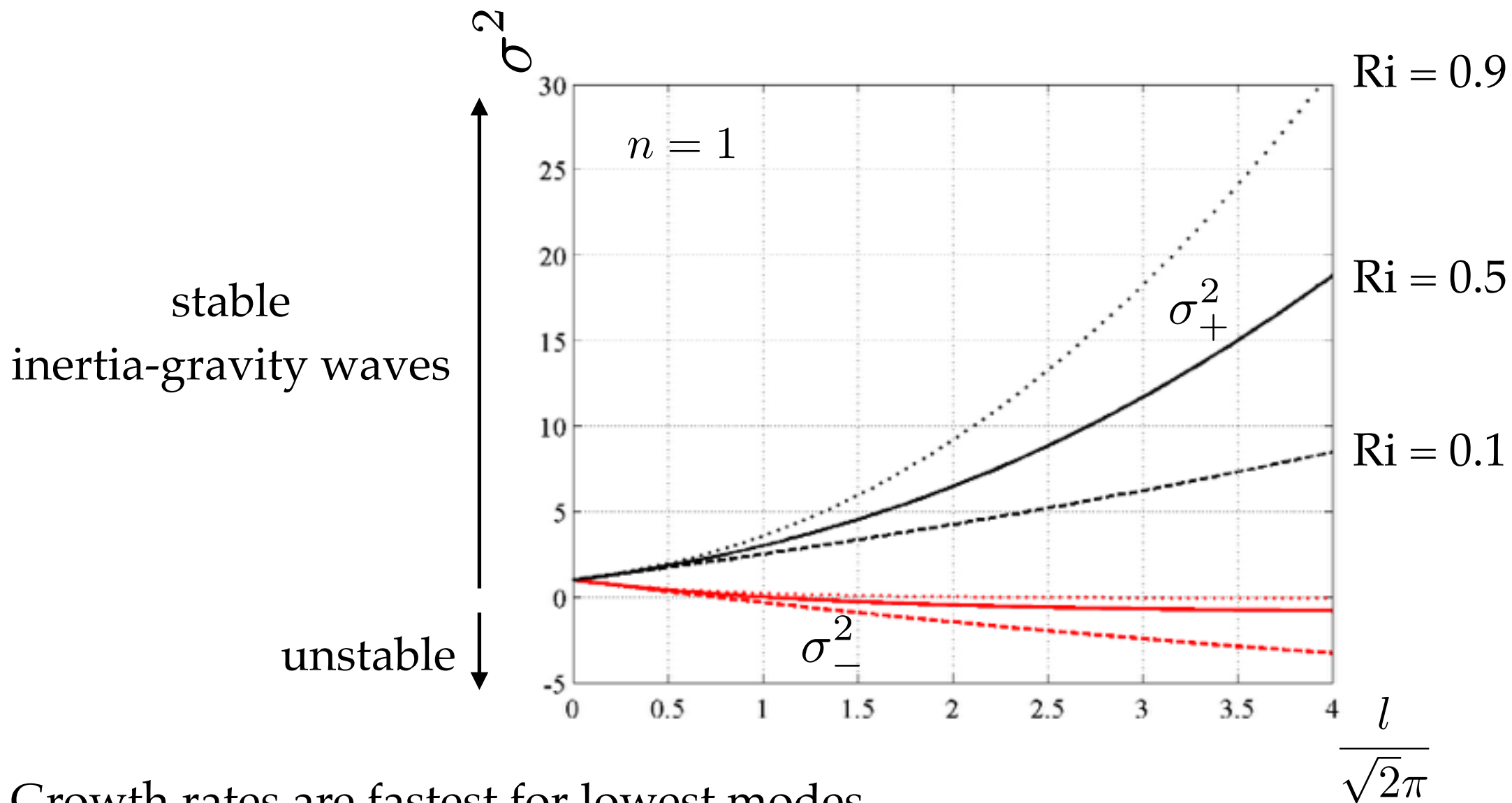
$\sigma_+^2$  is always stable

$\sigma_-^2$  is sometimes stable

Instability condition: 
$$\text{Ri} < 1 - \frac{n^2\pi^2}{l^2}$$

# There are stable and unstable solutions

$$\sigma^2 = 1 + \frac{\text{Ri}l^2}{2n^2\pi^2} \pm \sqrt{\frac{\text{Ri}^2l^4}{4n^4\pi^4} + \frac{l^2}{n^2\pi^2}}$$

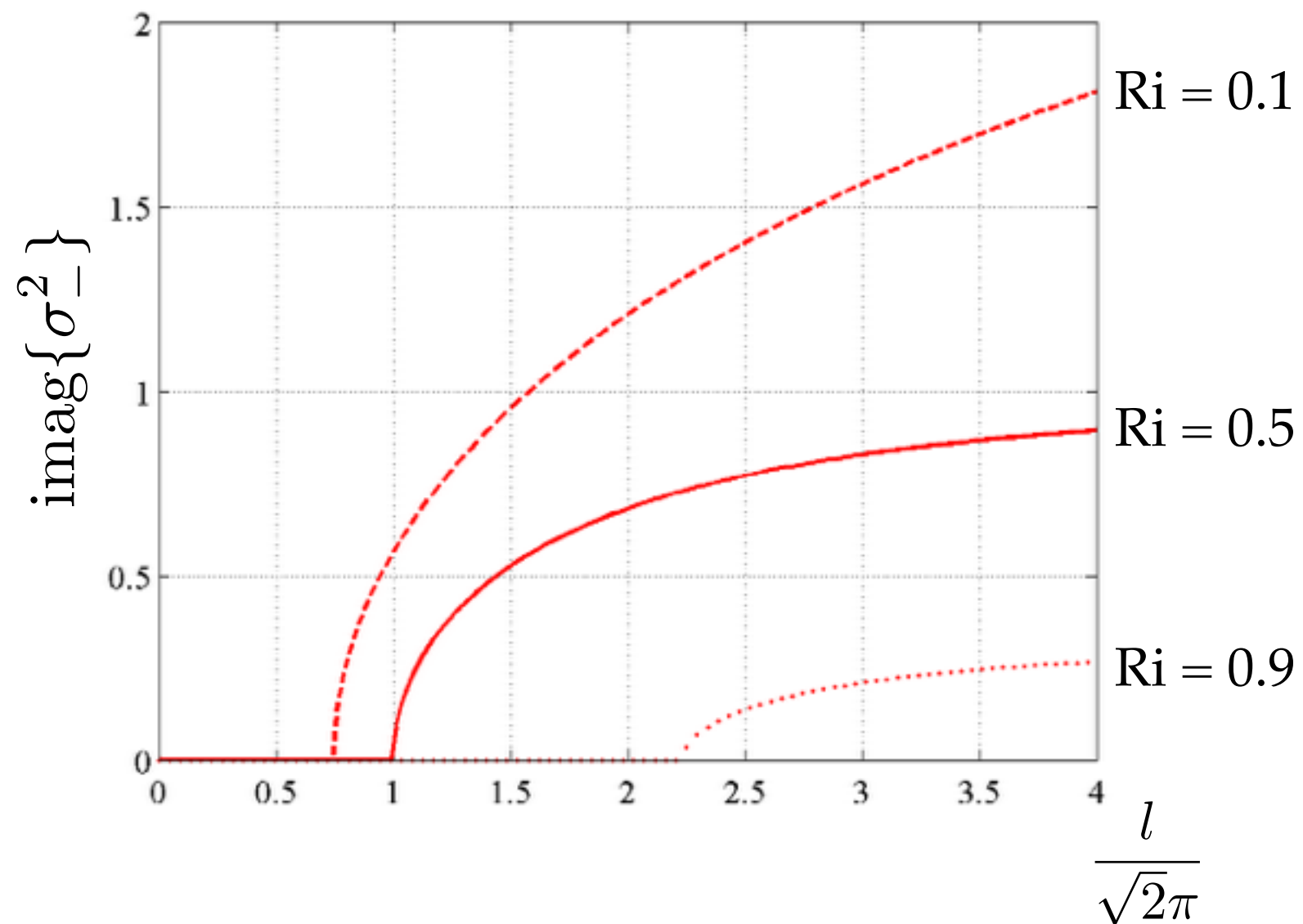


Growth rates are fastest for lowest modes



# Symmetric instability: growth rates

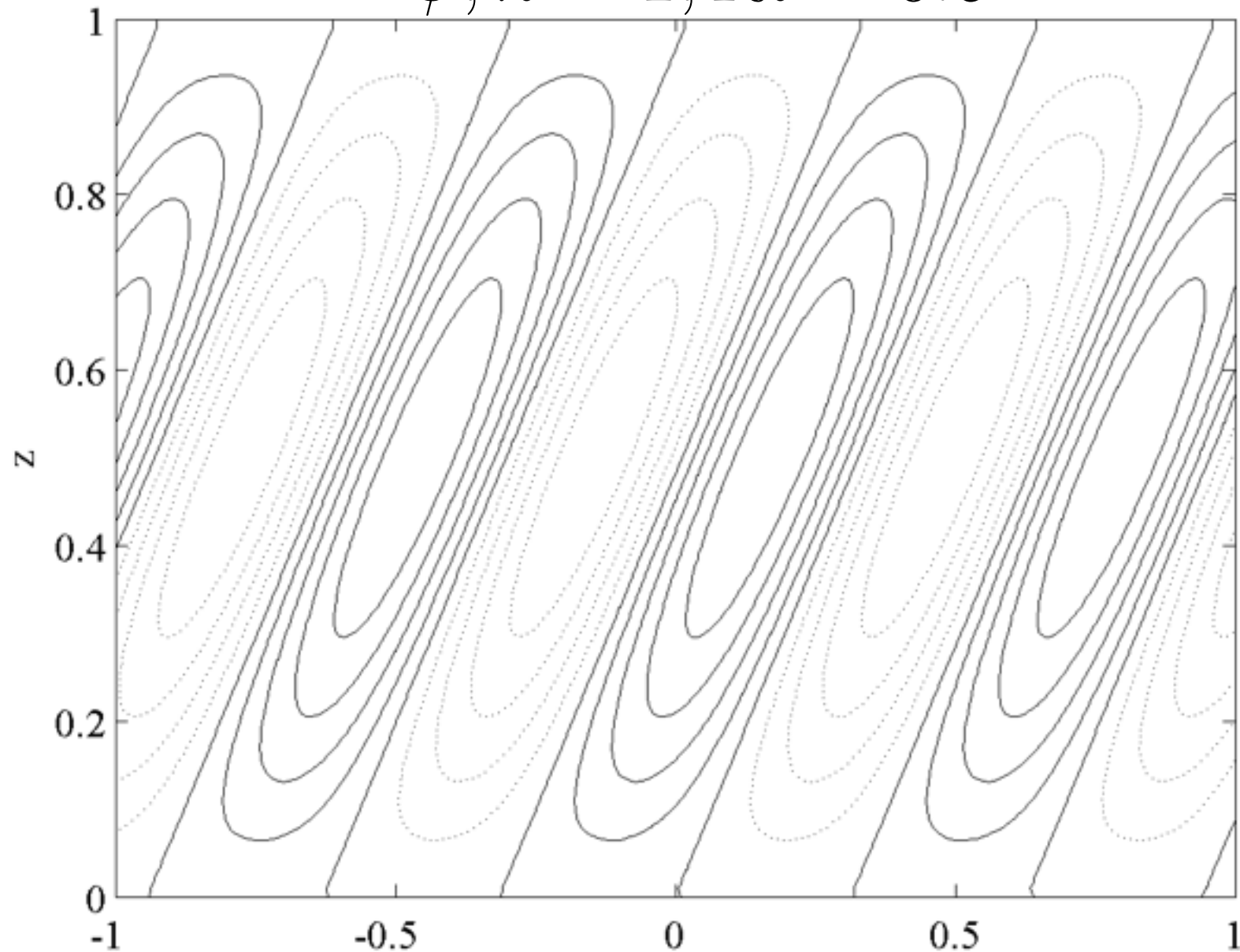
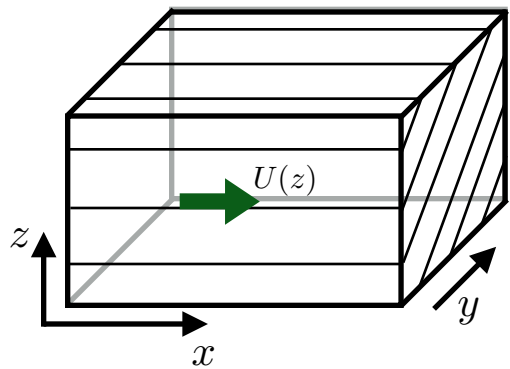
$$\sigma_-^2 = 1 + \frac{\text{Ri}l^2}{2n^2\pi^2} - \sqrt{\frac{\text{Ri}^2l^4}{4n^4\pi^4} + \frac{l^2}{n^2\pi^2}}$$



# Symmetric instability: eigenmodes

$$\psi = \text{Re}[(e^{im_+z} - e^{im_-z})e^{il_0y}]$$

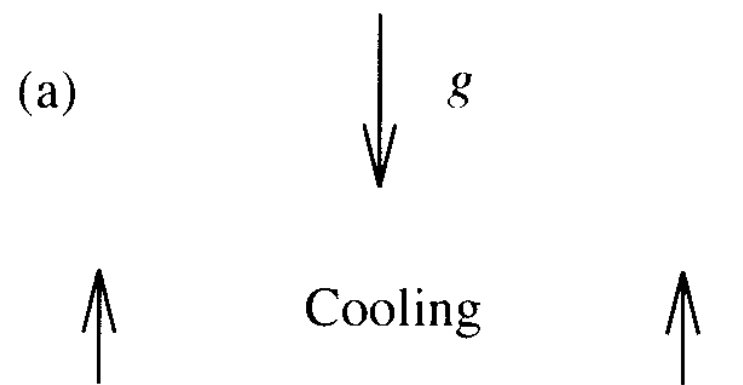
$$\psi, n = 1, Ri = 0.5$$



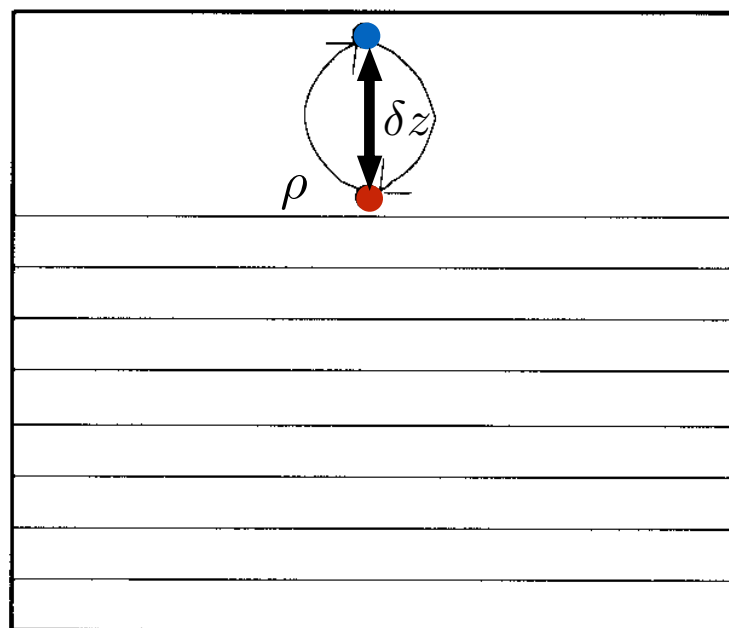
Instability forms as rolls aligned with the background flow

# Symmetric instability: a mixed gravitational-centrifugal instability

## 1. Gravitational instability: source is PE



$$\begin{aligned} \frac{Dw}{Dt} &= \frac{D^2}{Dt^2}(\delta z) = -g - \frac{1}{\rho} \frac{\partial p}{\partial z} \\ &= g \left( \frac{\rho_0 - \rho}{\rho} \right) \\ &= -N^2 \delta z \end{aligned}$$



$$\frac{\partial p_0}{\partial z} = -\rho_0 g$$

$$N^2 > 0 \text{ stable}$$

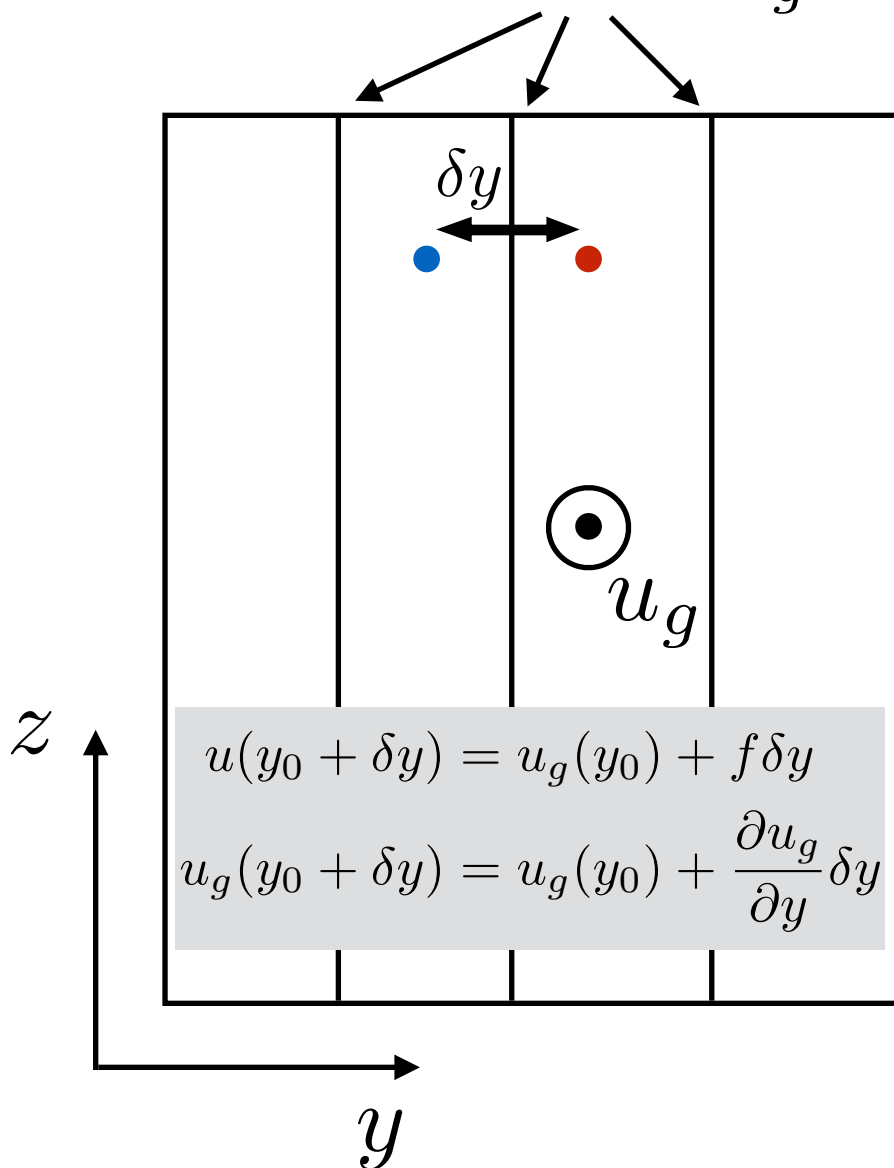
$$N^2 = 0 \text{ neutral}$$

$$N^2 < 0 \text{ unstable}$$

# Symmetric instability: a mixed gravitational-centrifugal instability

## 2. Centrifugal instability: source is KE

Lines of constant  $M$  for  $u_g = u_g(y)$



$$\begin{aligned} \frac{Dv}{Dt} &= \frac{D^2}{Dt^2}(\delta y) = f(u_g - u) \\ &= -f \left( f - \frac{\partial u_g}{\partial y} \right) \delta y \\ &= -f \frac{\partial M}{\partial y} \delta y \end{aligned}$$

Absolute momentum:  $M = fy - u_g$

$$f \frac{\partial M}{\partial y} > 0 \text{ stable}$$

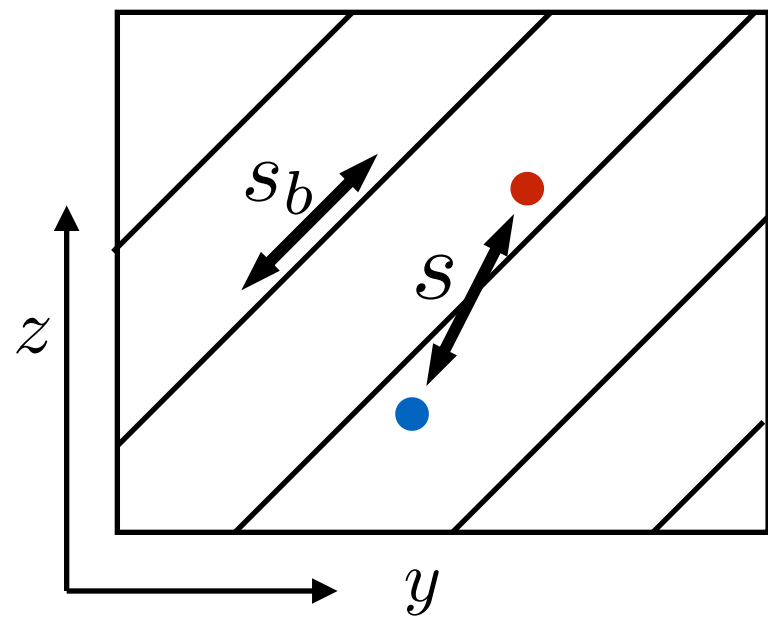
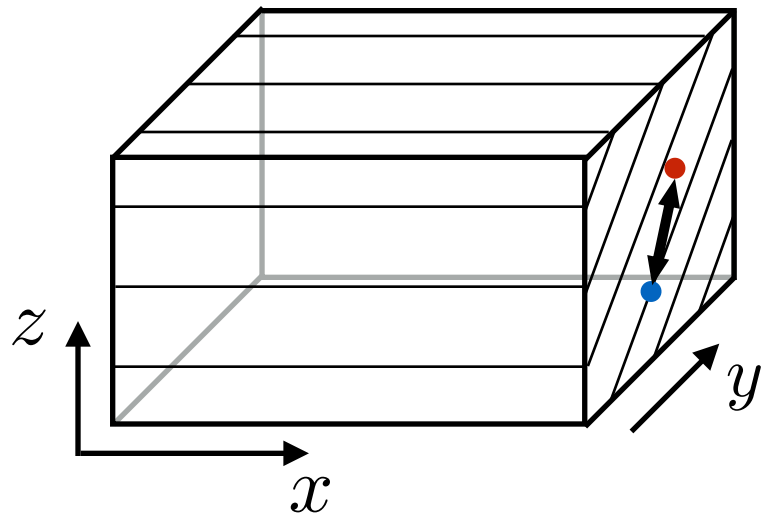
$$f \frac{\partial M}{\partial y} = 0 \text{ neutral}$$

$$f \frac{\partial M}{\partial y} < 0 \text{ unstable}$$

*following Holton*

# Symmetric instability: a mixed gravitational-centrifugal instability

## 3. Symmetric instability

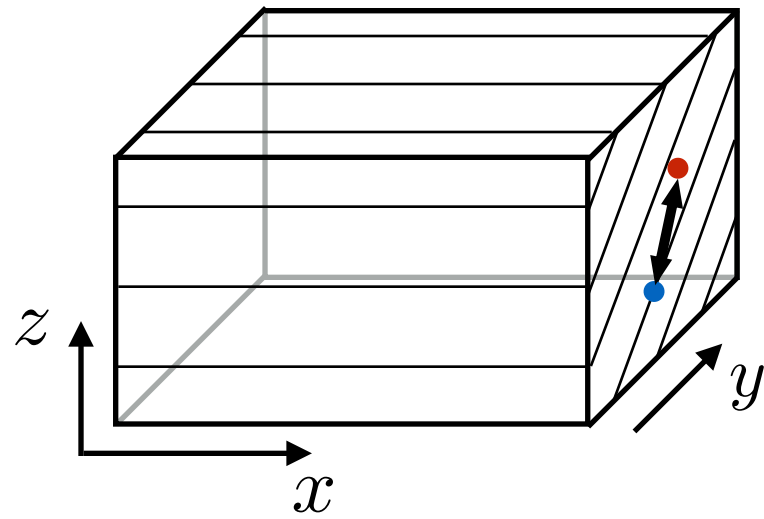


$$s = \frac{z_2 - z_1}{y_2 - y_1}$$

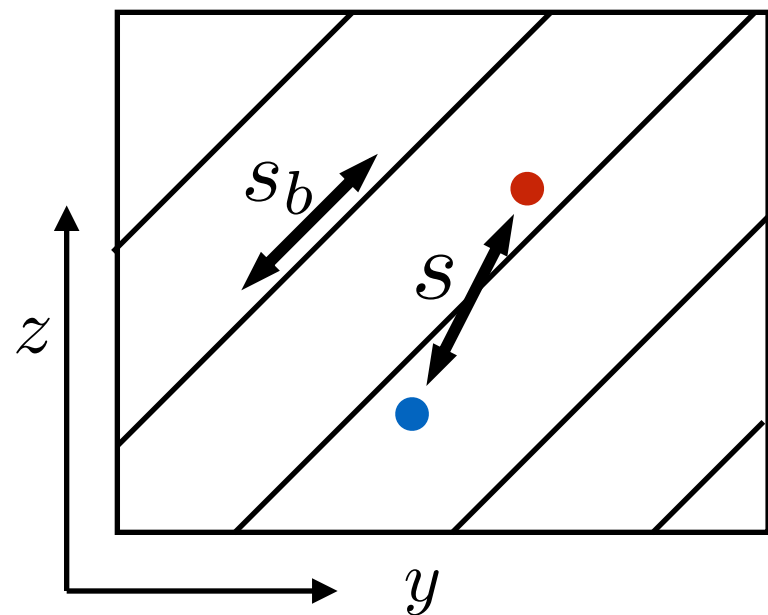
$$s_b = -\frac{M^2}{N^2}$$

# Symmetric instability: a mixed gravitational-centrifugal instability

## 3. Symmetric instability



$$\begin{aligned}\Delta PE &= \rho_0(z_2 - z_1)(b_2 - b_1) \\ &= \rho_0(z_2 - z_1)[M^2(y_2 - y_1) + N^2(z_2 - z_1)] \\ &= \rho_0 N^2 \Delta y^2 s(s - s_b)\end{aligned}$$



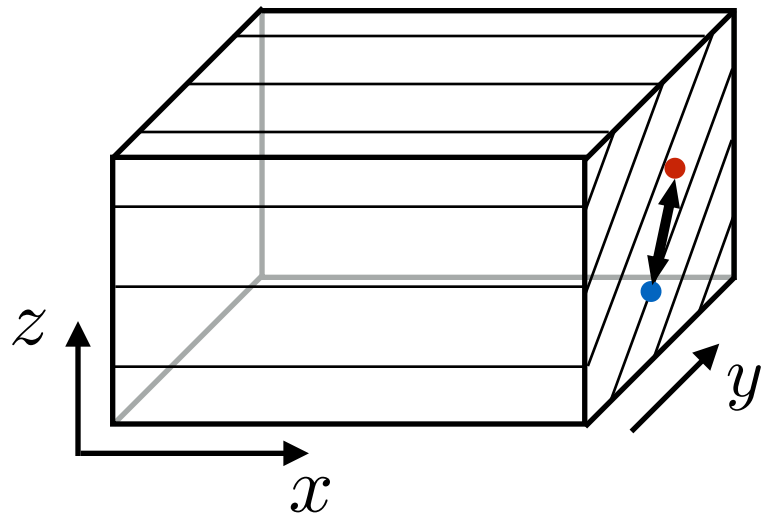
$$s = \frac{z_2 - z_1}{y_2 - y_1}$$

$$s_b = -\frac{M^2}{N^2}$$

$$\delta b = \bar{b}_y \delta y + \bar{b}_z \delta z$$

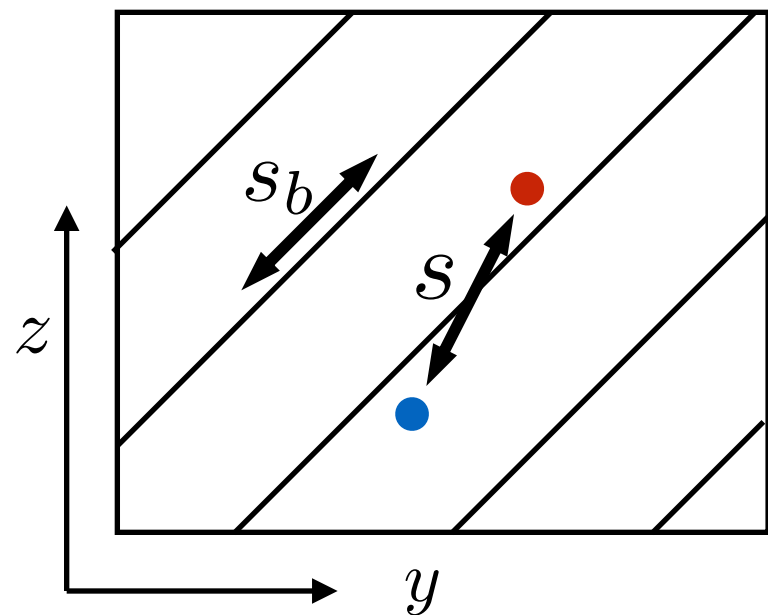
# Symmetric instability: a mixed gravitational-centrifugal instability

## 3. Symmetric instability



$$\begin{aligned}\Delta PE &= \rho_0(z_2 - z_1)(b_2 - b_1) \\ &= \rho_0(z_2 - z_1)[M^2(y_2 - y_1) + N^2(z_2 - z_1)] \\ &= \rho_0 N^2 \Delta y^2 s(s - s_b)\end{aligned}$$

$$\begin{aligned}\Delta KE &= (1/2)\rho_0[\{u_1 + f(y_2 - y_1)\}^2 + \\ &\quad \{u_2 - f(y_2 - y_1)\}^2 - u_1^2 - u_2^2] \\ &= \rho_0(y_2 - y_1)^2 f \left( f - s \frac{\partial u}{\partial z} \right) \\ &= \rho_0 \Delta y^2 [f^2 - N^2 s s_b]\end{aligned}$$

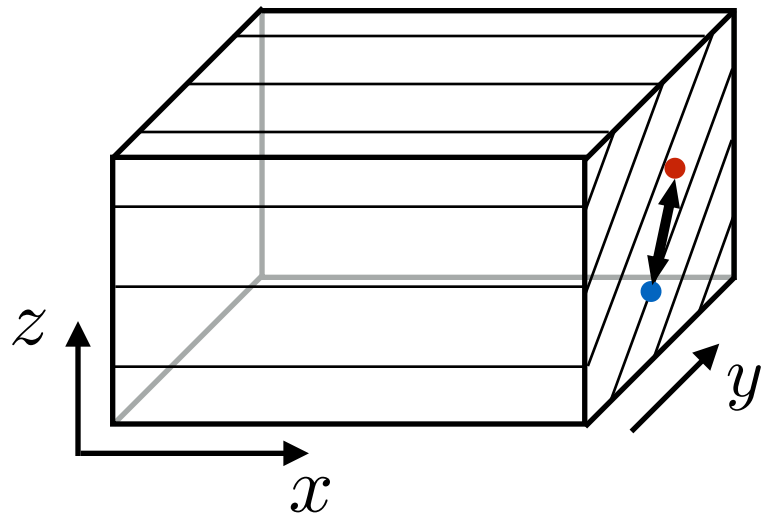


$$s = \frac{z_2 - z_1}{y_2 - y_1}$$

$$s_b = -\frac{M^2}{N^2}$$

# Symmetric instability: a mixed gravitational-centrifugal instability

## 3. Symmetric instability



$$\begin{aligned}\Delta PE &= \rho_0(z_2 - z_1)(b_2 - b_1) \\ &= \rho_0(z_2 - z_1)[M^2(y_2 - y_1) + N^2(z_2 - z_1)] \\ &= \rho_0 N^2 \Delta y^2 s(s - s_b)\end{aligned}$$

$$\begin{aligned}\Delta KE &= (1/2)\rho_0[\{u_1 + f(y_2 - y_1)\}^2 + \\ &\quad \{u_2 - f(y_2 - y_1)\}^2 - u_1^2 - u_2^2]\end{aligned}$$

$$= \rho_0(y_2 - y_1)^2 f \left( f - s \frac{\partial u}{\partial z} \right)$$

$$= \rho_0 \Delta y^2 [f^2 - N^2 s s_b]$$

$$s = \frac{z_2 - z_1}{y_2 - y_1}$$

$$s_b = -\frac{M^2}{N^2}$$

$$\Delta(KE + PE) = \rho_0 \Delta y^2 [f^2(1 - 1/Ri) + N^2(s - s_b)^2]$$

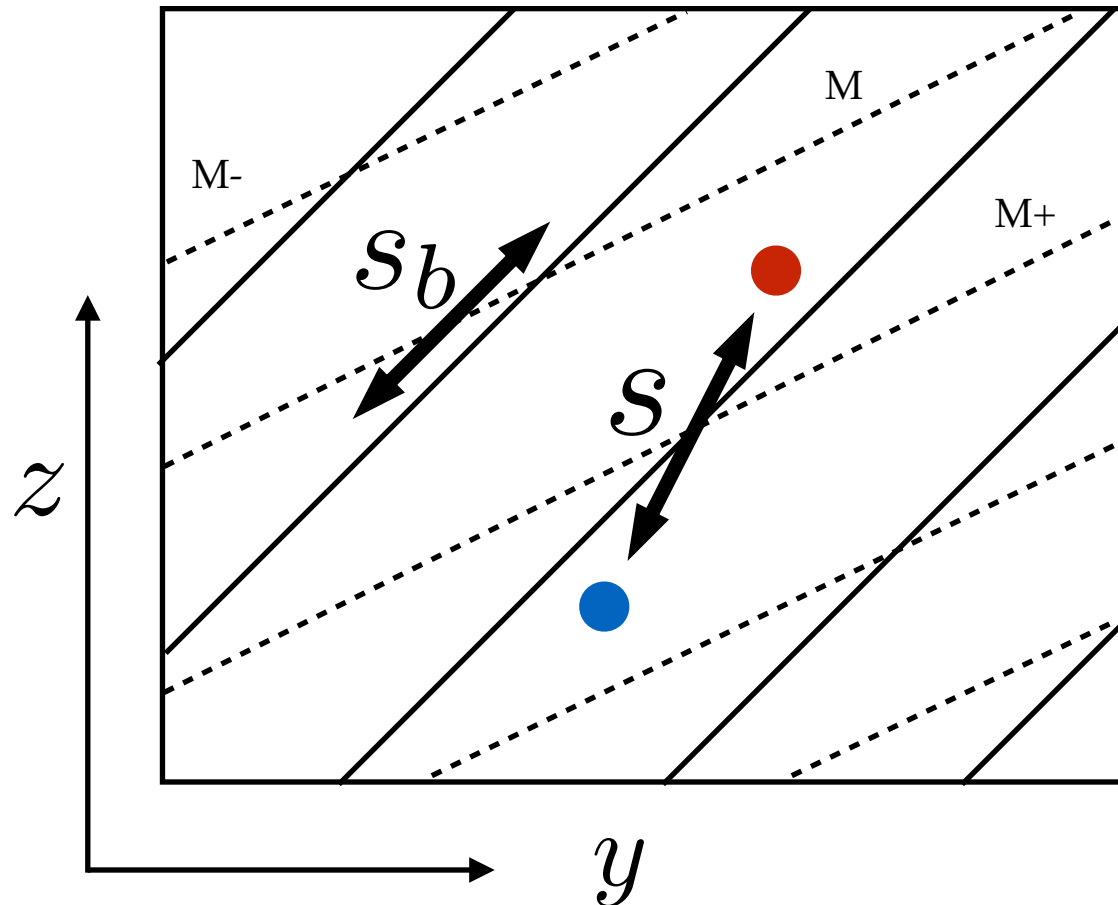
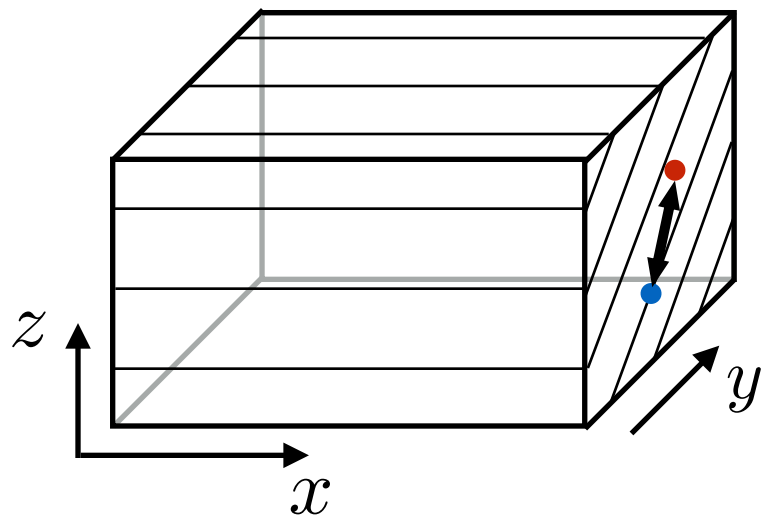
For instability, energy change must be -ve, so  $Ri < 1$

*Haine & Marshall, JPO 1998*



# Symmetric instability: a mixed gravitational-centrifugal instability

## 3. Symmetric instability



- Slopes of absolute momentum can be less than isopycnal slopes in regions of weak vertical stratification and strong horizontal stratification.
- Also known as “isentropic inertial instability”.

# Non-geostrophic baroclinic instability is qualitatively similar to Eady instability: $l=0$

$$[1 - (\sigma + kU)^2] \partial_{zz} w - \frac{2k}{\sigma + kU} \partial_z w - Rik^2 w = 0$$

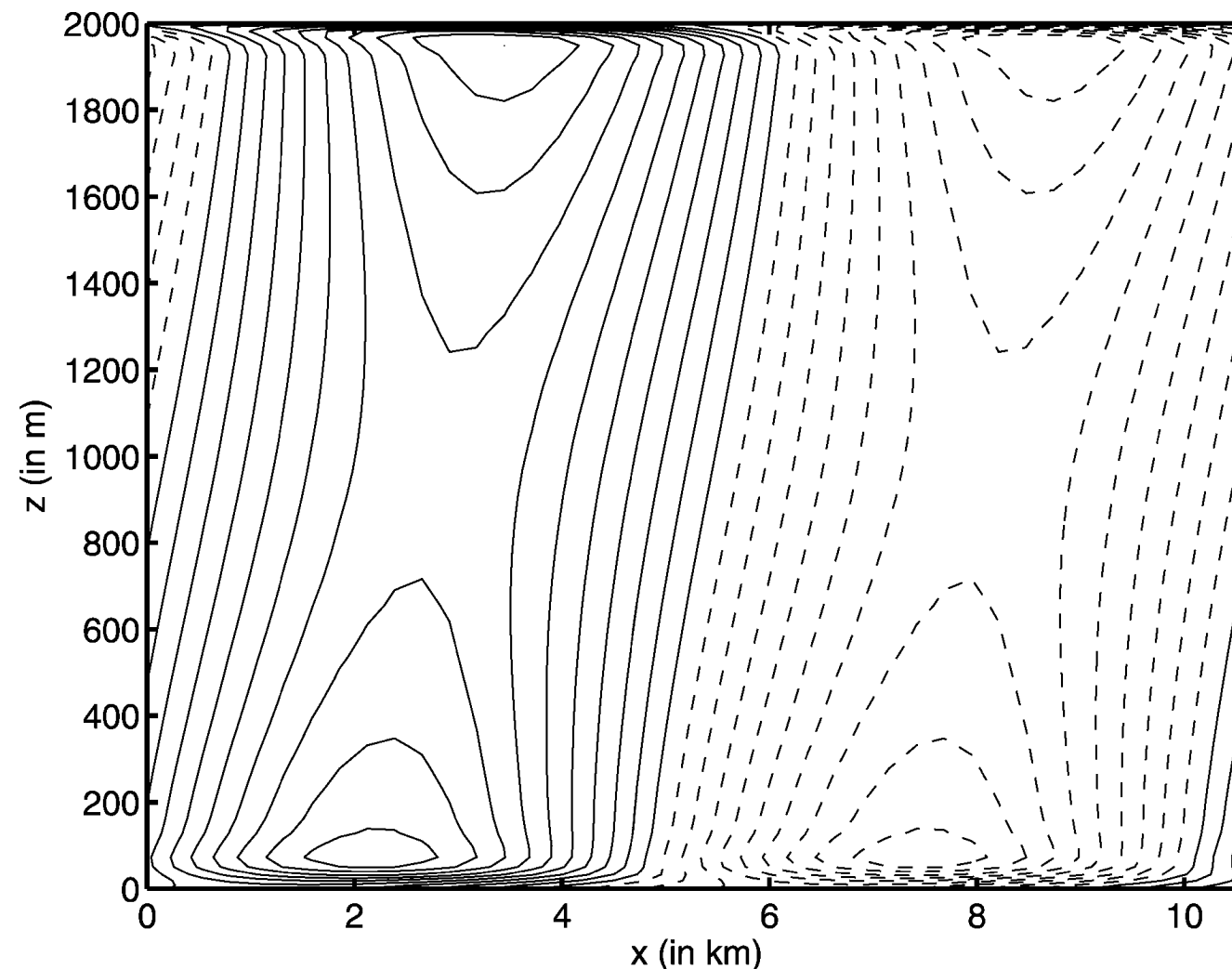


FIG. 4. A snapshot of the temperature perturbation for the fastest-growing mode.

Unstable for  $Ri > 1$

# Stone: Non-geostrophic stability with perturbations in 2 directions

$$[1 - (\sigma + kU(z))^2]w'' - 2 \left[ \frac{k}{\sigma + kU(z)} - il \right] w' - \left[ Ri(k^2 + l^2) - \frac{2ikl}{\sigma + kU(z)} \right] w = 0$$

Boundary conditions  $w(0) = w(1) = 0$

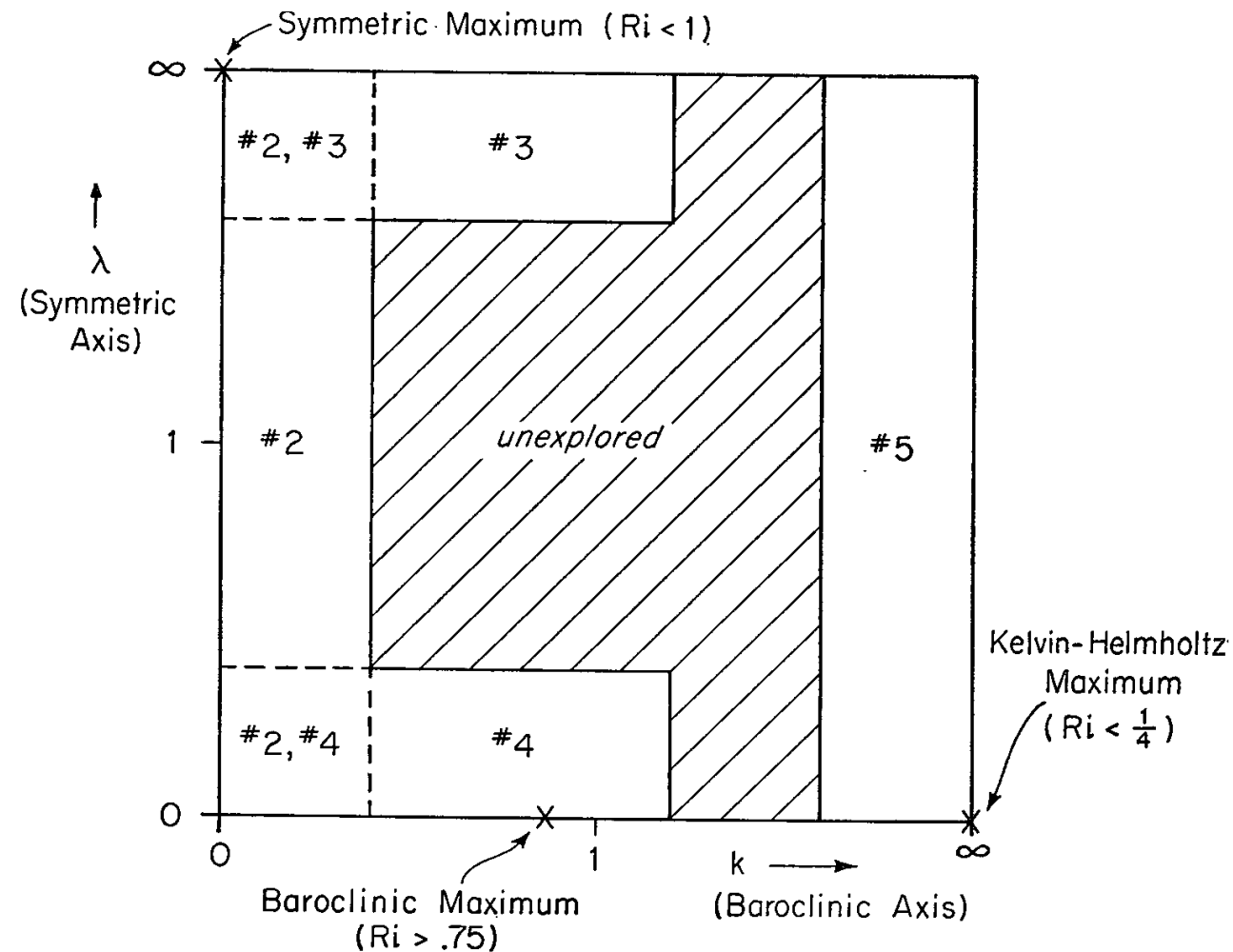
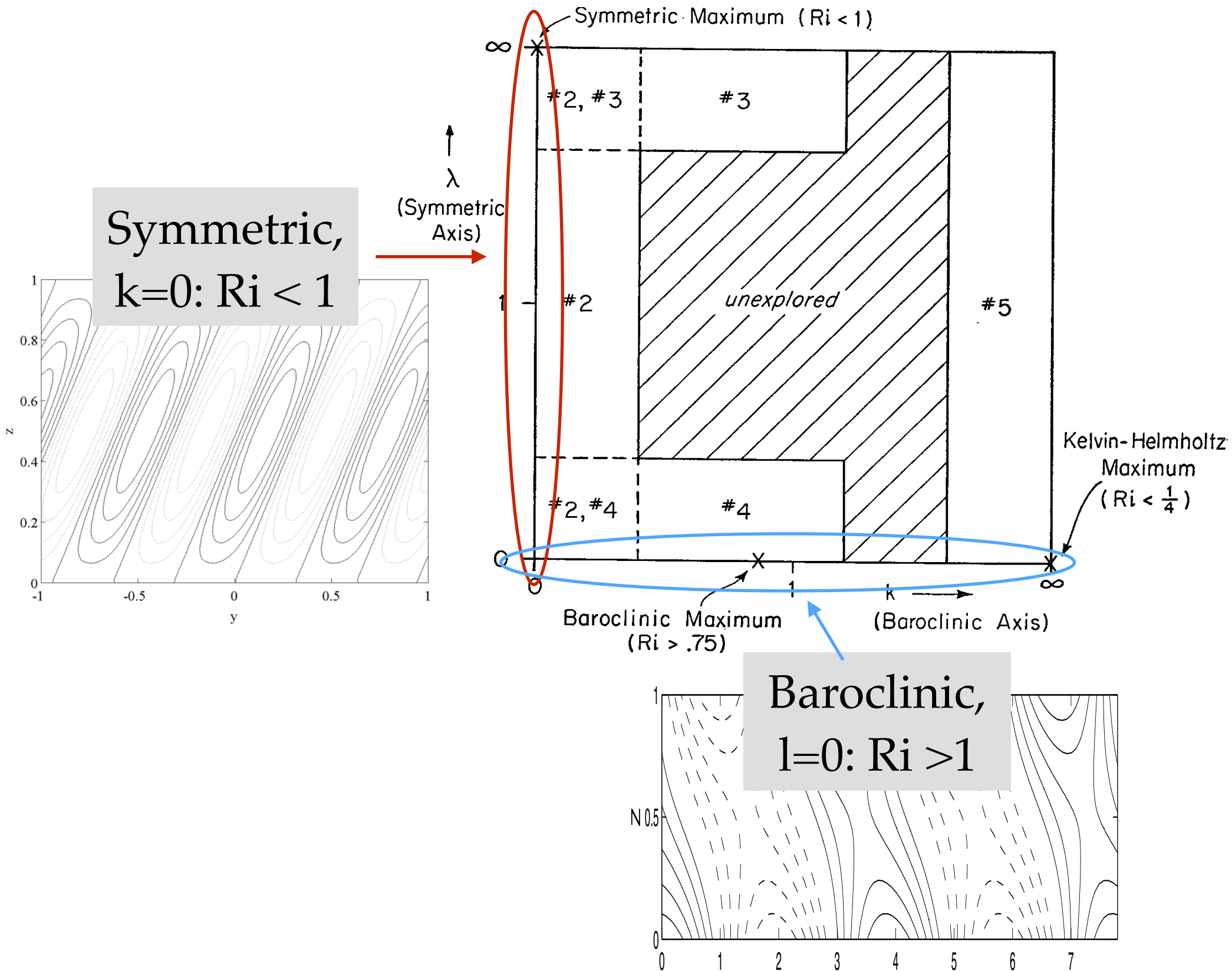


FIG. 5. Schematic diagram of the  $k$ - $\lambda$  plane, showing the various regions where the analyses of Sections 2 through 5 apply. The locations of the three local maxima in the growth rate are indicated schematically by  $x$ 's, and the range of values of  $Ri$  for which these maxima exist are also given.

# Stone: Non-geostrophic stability with perturbations in 2 directions



# Non-geostrophic instabilities arising from 2D perturbations

Ri = 2:

~Baroclinic. Max growth rate at  $l=0$ .  
Single unstable mode.

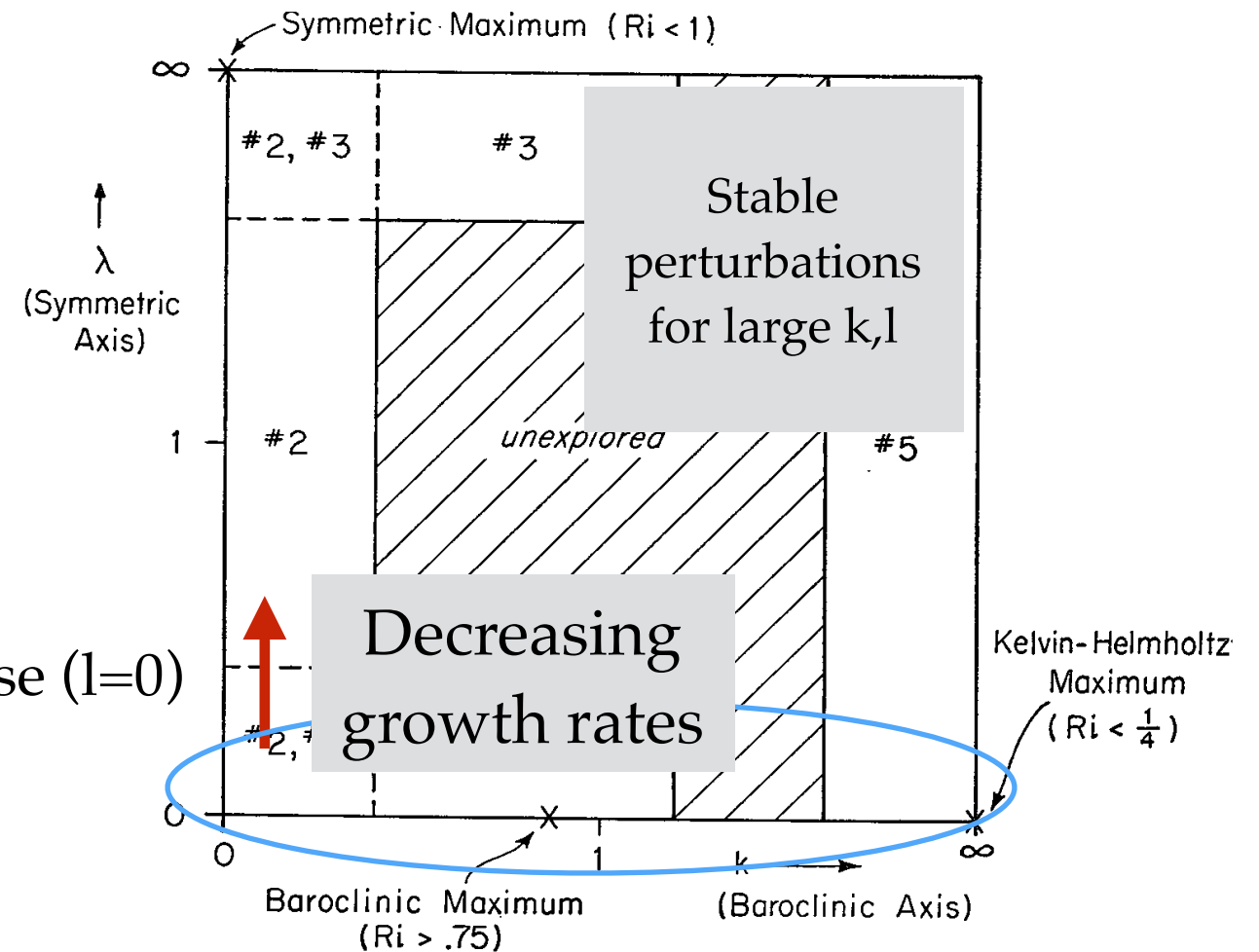
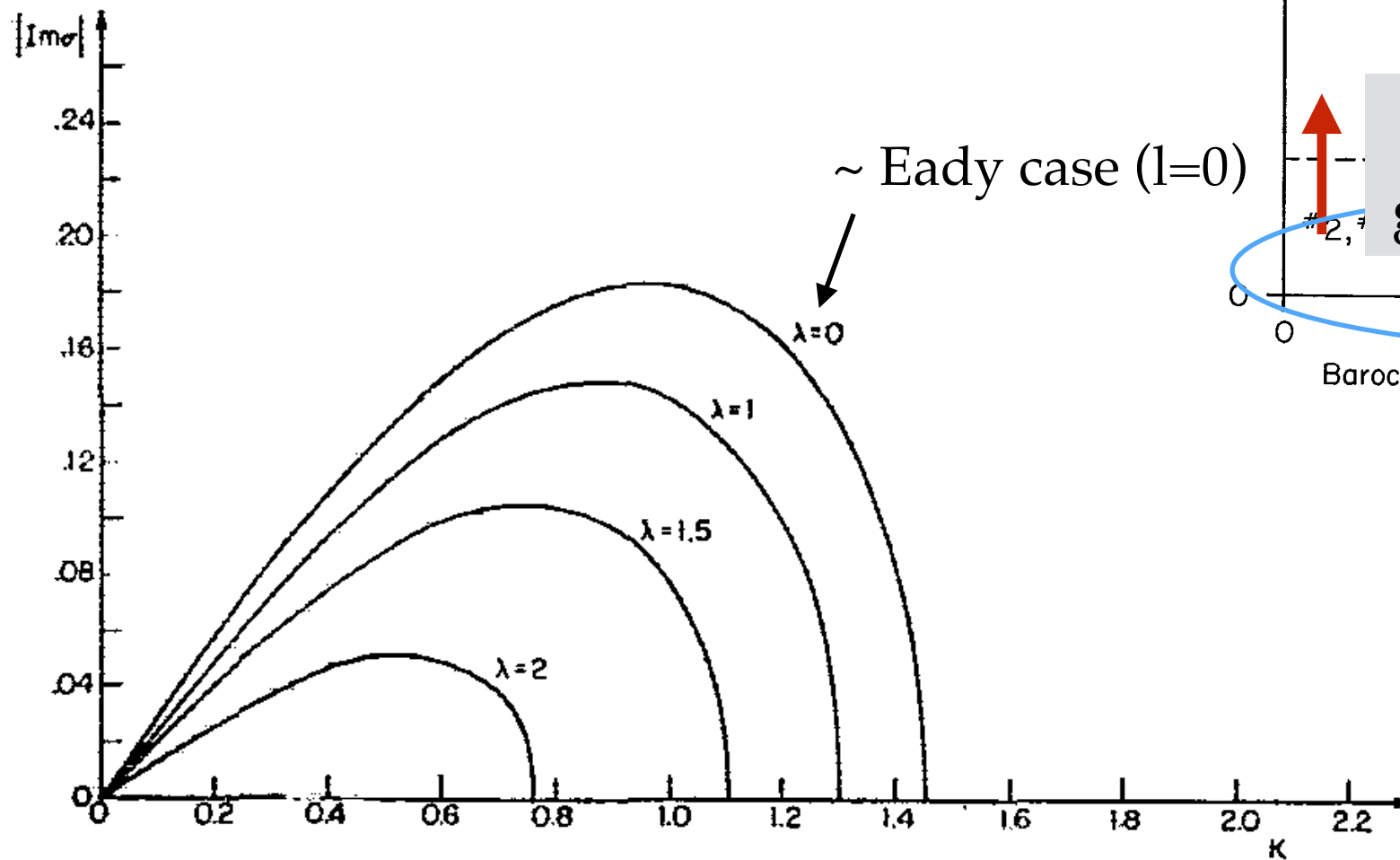
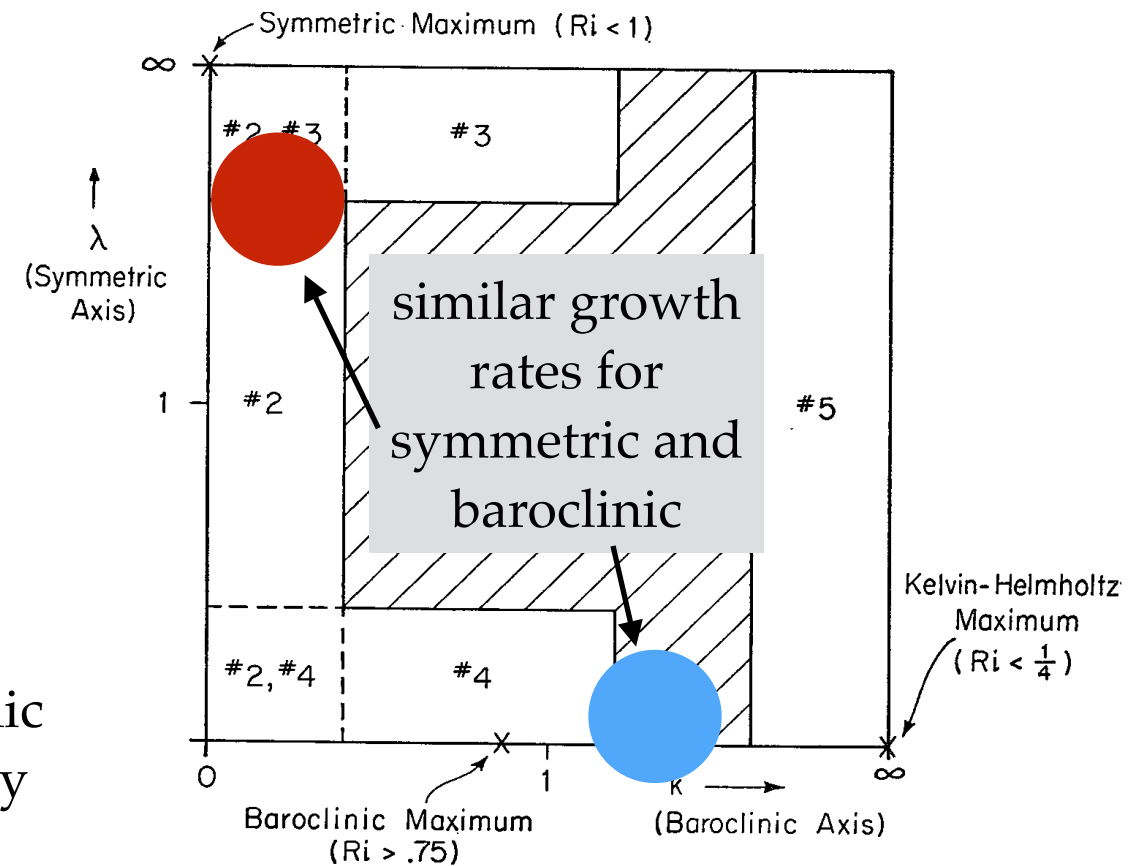
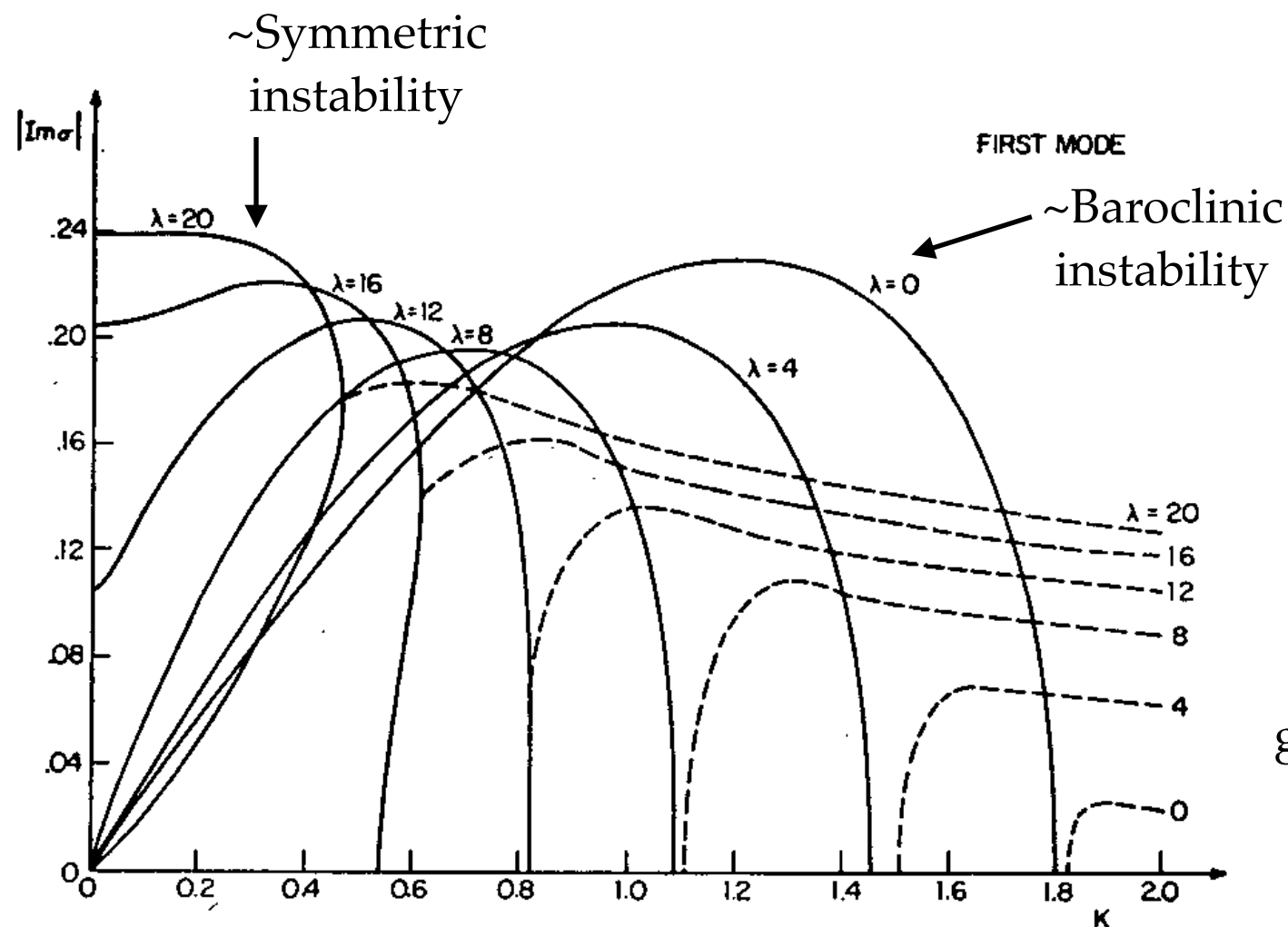


FIG. 1. Growth rates vs  $k$  when  $Ri=2$ .

# Non-geostrophic instabilities arising from 2D perturbations

**Ri = 0.92:**

More than one vertical mode unstable. Symmetric instability appears.



Secondary unstable modes have smaller growth rates and smaller scales

FIG. 3. Growth rates for the most unstable mode when  $Ri=0.92$ .

# Non-geostrophic instabilities arising from 2D perturbations

**Ri = 0.5:**  
Symmetric has largest growth rate

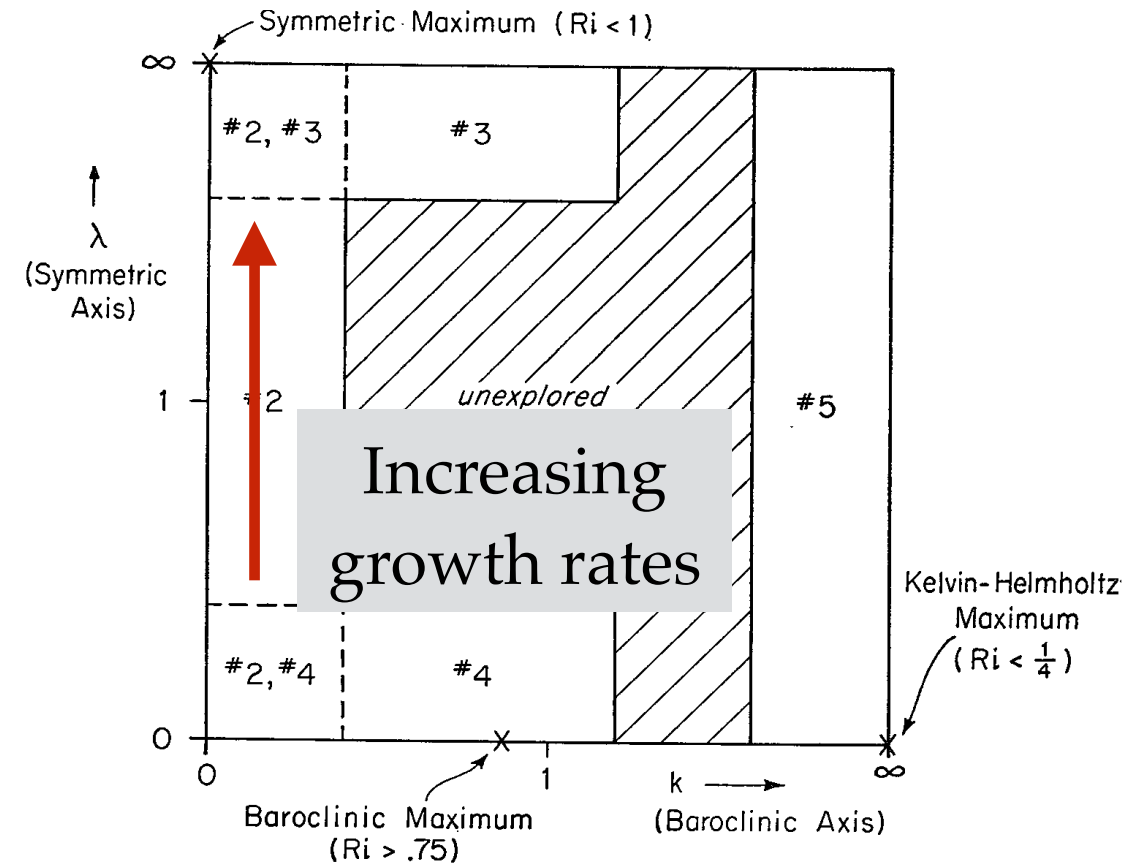
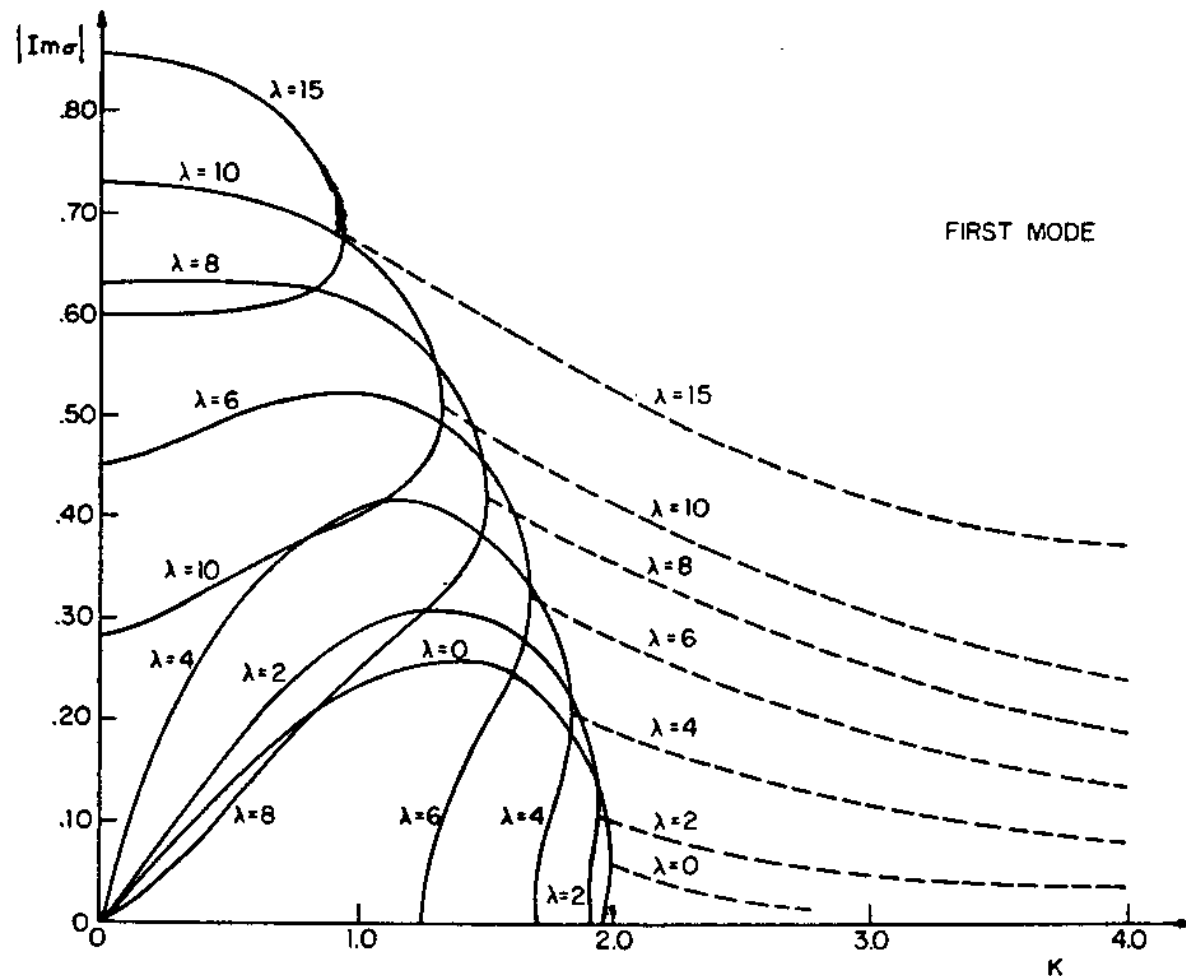


FIG. 4. Growth rates for the most unstable mode when  $Ri=0.5$ .

# Summary

- A flow in thermal wind balance is subject to a range of instabilities when the full primitive equations are considered
  - $0 < Ri < 0.25 \rightarrow$  Kelvin Helmholtz
  - $0.25 < Ri < 0.95 \rightarrow$  Symmetric
  - $0.95 < Ri \rightarrow$  Baroclinic
- Baroclinic instability is not significantly modified in full equations compared to QG for  $Ri \gg 1$
- Symmetric instability can be thought of as a combined centrifugal and gravitational instability